

# Interactions between sediment microbial ecology and physical dynamics drive heterogeneity in contextually similar depositional systems

Hope, Julie A.; Malarkey, Jonathan; Baas, Jaco H.; Peakall, Jeff; Parsons, Daniel; Manning, Andrew J.; Bass, Sarah; Lichtman, Ian; Thorne, Peter; Ye, Leiping; Paterson, David M.

### Limnology and Oceanography

DOI: https://doi.org/10.1002/lno.11461

Published: 01/10/2020

Peer reviewed version

Cyswllt i'r cyhoeddiad / Link to publication

*Dyfyniad o'r fersiwn a gyhoeddwyd / Citation for published version (APA):* Hope, J. A., Malarkey, J., Baas, J. H., Peakall, J., Parsons, D., Manning, A. J., Bass, S., Lichtman, I., Thorne, P., Ye, L., & Paterson, D. M. (2020). Interactions between sediment microbial ecology and physical dynamics drive heterogeneity in contextually similar depositional systems. *Limnology and Oceanography*, *65*(10), 2403-2419. https://doi.org/10.1002/lno.11461

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Limnology and Oceanography



## Interactions between sediment microbial ecology and physical dynamics drive heterogeneity in contextually similar depositional systems.

Journal:	Limnology and Oceanography				
Manuscript ID	LO-19-0429.R1				
Wiley - Manuscript type:	Original Article				
Date Submitted by the Author:	n/a				
Complete List of Authors:	Hope, Julie; The University of Auckland, Institute of Marine Science; University of Saint Andrews, Scottish Oceans Institute Malarkey, Jonathan; Bangor University, School of Ocean Sciences; University of Hull, Department of Geography, Environment and Earth Sciences Baas, Jaco; Bangor University, School of Ocean Sciences Peakall, Jeffrey; University of Leeds, School of Earth and Environment Parsons, Daniel ; University of Hull, Department of Geography, Environment and Earth Sciences Manning, Andrew; HR Wallingford Ltd; University of Hull, Department of Geography, Environment and Earth Sciences; University of Plymouth, School of Marine Science and Engineering; University of Delaware, Department of Civil and Environmental Engineering, Center for Applied Coastal Research Bass, Sarah ; University of Plymouth, School of Marine Science and Engineering Lichtman, Ian ; National Oceanography Centre Liverpool Thorne, Peter; National Oceanography Centre Liverpool Ye, Leiping; University of Hull, Department of Geography, Environment and Earth Sciences; University of Delaware, Department of civil and environmental engineering Center for Applied Coastal Research				
Keywords:	microphytobenthos, biostabilisation, temporal dynamics, biofilm,				
Keywords.	sediment erosion				
Abstract:	This study focuses on the strong interactions between the stability of different sediments and the biological and physical variables that influence the erodibility of the bed. Sampling at short-term temporal scales illustrated the persistence of the microphytobenthic (MPB) community even during periods of frequent, high physical disturbance. The role of MPB in biological stabilisation along a changing sedimentary habitat was also assessed . Key biological and physical properties, such as the MPB biomass, composition and extracellular polymeric substances, were used to predict sediment stability (erosion threshold) of muddy and sandy habitats within close proximity to one another over multiple days as well as within emersion periods. This allowed the effects of dewatering, MPB growth and productivity to be examined as well as				

the resilience and recovery of the MPB community after physical disturbance from tidal currents and wave exposure.
Canonical analysis of principal components (CAP) ordinations were used to illustrate the trends observed in bio-physical properties between the sites, while marginal and sequential distance-based linear models (DistLM) were used to identify key properties influencing sediment erodibility. While grain size was important for site differences in the CAP analysis, it contributed less to the variability in sediment erodibility than other key biological parameters. Among the biological predictors, MPB diversity explained very little variation in marginal tests but was a significant predictor in sequential tests when MPB biomass was also considered with diversity and biomass key predictors of sediment stability, contributing 9% and 10% respectively to the final model across all sites.

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- 2 contextually similar depositional systems.
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- 4 Hope, JA<sup>1,2</sup>, Malarkey, J<sup>3,4</sup>, Baas, JH<sup>3</sup>, Peakall, J<sup>5</sup>, Parsons, DR<sup>4</sup>, Manning, AJ<sup>4,6,7,9</sup>, Bass, SJ<sup>6</sup>,
- 5 Lichtman, ID<sup>8</sup>, Thorne, PD<sup>8</sup>, Ye, L<sup>4,9</sup> and Paterson, DM<sup>1</sup>
- 6 <sup>1</sup>Sediment Ecology Research Group, School of Biology, University of St Andrews, Scotland
- 7 <sup>2</sup> Institute of Marine Science, University of Auckland, New Zealand
- 8 <sup>3</sup> School of Ocean Sciences, Bangor University, Menai Bridge, Isle of Anglesey, Wales.
- 9 <sup>4</sup> Energy & Environment Institute, University of Hull, England
- 10 <sup>5</sup> School of Earth and Environment, University of Leeds, England
- 11<sup>6</sup> School of Biological and Marine Sciences, University of Plymouth, England
- <sup>7</sup> HR Wallingford, Howbery Park, England
- 13 <sup>8</sup>National Oceanography Centre, Joseph Proudman Building, 6 Brownlow Street, Liverpool, England
- <sup>9</sup> Department of Civil and Environmental Engineering, Center for Applied Coastal Research,
- 15 University of Delaware, USA
- 16 Correspondence to: julie.hope@auckland.ac.nz
- 17
- 18 Keywords: Biological stabilisation, biostabilisation, estuaries, ecosystem function, erodibility,
- 19 extracellular polymeric substances, EPS, microphytobenthos, MPB, sediment stability, short-term
- 20 temporal dynamics, sediment dynamics.
- 21 Running head: Bio-physical mediation of sediment dynamics

### 22 Abstract

23 This study focuses on the interactions between the stability of different sediments and the biological 24 and physical variables that influence the erodibility of the bed. Sampling at short-term temporal scales illustrated the persistence of the microphytobenthic (MPB) community even during periods of frequent, 25 26 high physical disturbance. The role of MPB in biological stabilisation of the sediment along a changing sedimentary habitat was also assessed. Key biological and physical properties, such as the MPB 27 biomass, composition and extracellular polymeric substances (EPS), were used to predict the sediment 28 29 stability (erosion threshold) of muddy and sandy habitats within close proximity to one another over 30 multiple days, as well as within emersion periods. This allowed the effects of dewatering, MPB growth 31 and productivity to be examined as well as the resilience and recovery of the MPB community after 32 physical disturbance caused by tidal currents and waves.

33 Canonical analysis of principal components (CAP) ordinations were used to visualise and assess the 34 trends observed in bio-physical properties between the sites, and marginal and sequential distance-based 35 linear models (DistLM) were used to identify the key properties influencing sediment erodibility. While 36 the particle size of the bed was important for differences between sites in the CAP analysis, it 37 contributed less to the variability in sediment erodibility than key biological parameters. Among the 38 biological predictors, MPB diversity explained very little variation in marginal tests but was a 39 significant predictor when MPB biomass was also considered in sequential tests. MPB diversity and 40 biomass were both key predictors of sediment stability, contributing 9% and 10% respectively to the 41 final model across all sites in comparison to 2% of the variance explained by sediment grain size.

42

43 Introduction

Variability in response to physical forcing is an inherent property of natural systems (Black et al., 2002)
and represents a significant challenge for modelling and predicting the behaviour of natural sediment
beds. It is also commonly suggested that an important source of this variability, biogenic stabilisation
(Parsons et al., 2016; Tolhurst et al., 2009), is largely confined to fine cohesive sediments (mud flats)
rather than more sandy substrata. However, this approach neglects the heterogeneous composition of
natural beds that vary both spatially and temporally (Rainey et al., 2003; Chapman et al., 2010).

50 Understanding the biogenic processes that generate variation and heterogeneity in natural systems will
 51 support our ability to model system behaviour more accurately in the future.

52 Biological stabilisation of cohesive (muddy) sediment is often attributed to microbially-produced 53 extracellular polymeric substances (EPS) that increase the cohesion between sediment particles, often 54 forming biofilms (Hubas et al., 2018). The production of EPS is typically attributed to bacteria and 55 microphytobenthos (MPB) (Chen et al., 2017; Lubarsky et al., 2010). While MPB are more abundant 56 on cohesive sediments, recent studies show that microbially-produced EPS can also hinder bedform 57 development and inhibit erosion in non-cohesive (sandy) (Chen et al., 2017; Malarkey et al., 2015) and 58 mixed sediment beds (Parsons et al., 2016). The balance between physical disturbance (hydrodynamic 59 stress) and bed erodibility is complex (Beninger et al., 2018). Regular physical forcing can restrict the 60 accumulation of fine sediment and MPB on the bed (Mariotti and Fagherazzi, 2012) preventing the 61 MPB standing stock from developing fully (Blanchard et al., 2001). However, once developed, after a period of calm conditions, increasing erosive stress may be resisted. The biomass and nature of the 62 63 MPB community, epipelic (e.g. Paterson and Hagerthey, 2001) or epipsammic (e.g. Harper and Harper, 64 1967; Hickman and Round, 2007), will contribute to the variability of the response to stress.

65 If a system is largely abiotic then the introduction of biota can create greater heterogeneity or 66 homogenise the system. For example, greater microphytobenthic diversity has been linked to higher 67 grazer diversity (Balvanera et al., 2006) which through differences in bioturbation can increase habitat 68 heterogeneity (Hale et al., 2015). Furthermore, patchy biofilm distribution and growth have been 69 associated with positive bio-physical feedbacks as the system becomes inherently more patchy leading 70 to spatial self-organisation and more fine sediment accretion and eventually influencing large 71 geomorphological features (Weerman et al., 2010). However, the introduction of biota can also cause 72 different sediments to become more similar to one another. For instance, the presence of large infauna can also 'smooth out' the effects of flow on sediment resuspension across different sediment types (Li 73 74 et al., 2017). The former processes of increasing heterogeneity suggest that microbial growth and EPS 75 accumulation can not only have a localized effect, but if growth becomes extensive, biostabilisation is capable of influencing ecosystem functionality at various spatial and temporal scales (Orvain et al., 76 77 2012; Ubertini et al., 2015) and although variability may increase (Chapman et al., 2010) this can have

78 system-wide implications. Many studies have focussed on seasonal and inter-annual variability 79 (Montani et al., 2003; Wal et al., 2010). However, the mechanisms that drive changes to the structure 80 of the system (Van de Koppel et al., 2001) and resilience and recovery from disturbance can occur on 81 relatively short spatial (Spilmont et al., 2011) and temporal scales (Orvain et al., 2012). Furthermore, 82 the variability observed from small spatial scales over short timeframes in intertidal environments can 83 be of the same order of magnitude as both seasonal and annual variability (Seuront and Leterme, 2006). 84 The importance of considering temporal scales has been highlighted in previous soft sediment studies 85 (Hewitt et al., 2006; Tolhurst et al., 2005a). However, short-term temporal dynamics that may influence 86 EPS accumulation, biofilm development and biostabilisation have not been well characterised across 87 different sediment types. The development of biofilms depends on the balance between growth and 88 detachment, with hydrodynamic stress being a primary driver of benthic biofilm detachment (Telgmann 89 et al., 2004). We therefore require further information on the interactions between biofilm properties, biostabilisation, hydrodynamic stress and subsequent resistance to erosion over multiple emersion 90 91 periods and within different habitats. This information is essential to assess the role of biostabilisation, 92 both from ecological and dynamic perspectives (de Brouwer et al., 2000; Mariotti and Fagherazzi, 2012; 93 Underwood and Paterson, 2003). MPB influence on sediment stability and this key ecosystem function 94 augments their important roles in: The transfer of energy to higher organisms (MacIntyre et al., 1996); 95 the bentho-pelagic exchange of sediment (Chen et al., 2017); and nutrient cycling (McGlathery et al., 96 2004). MPB importance in these ecosystems highlights the need to understand the dynamics governing 97 their presence across different habitats. Frequent resuspension of MPB cells and related EPS may 98 prevent the formation of substantial biofilms, and therefore limit their biostabilisation potential (Aspden 99 et al., 2004), however, an "inoculum" often remains in place (Chen et al., 2017) leading to rapid re-100 colonisation under suitable conditions.

We hypothesise that biofilm properties such as MPB biomass, colloidal carbohydrate concentrations and the MPB diversity will influence biostabilisation of various sediment types. Furthermore, we hypothesise that the biogenic influence will persist over various temporal scales (emersion on consecutive days), as the microphytobenthic community and biotic characteristics tolerate regular, high intensity tidal inundation. As laboratory experiments cannot generally capture the natural variability in

large, complex and dynamic marine systems (Tolhurst et al., 2009) we examined these short-term 106 107 dynamics in closely associated muddy and sandy habitats in the Dee Estuary, England. This estuary is 108 subject to strong current velocities and frequent wave action, resulting in turbid waters with a high 109 suspended load (Amoudry et al., 2014), but has various sedimentary habitats in close proximity to one 110 another, making it an excellent model system. Suspended sediment often affects water quality, which limits light availability for sediment dwelling photosynthetic organisms during tidal inundation (Pratt 111 112 et al., 2014), and the physical disturbance from flow itself may prevent the accumulation of EPS and biofilm development on the bed (Blanchard et al., 1997; Ubertini et al., 2015). 113

114

115 Methods

*Study sites & sample collection* – The Dee is a hypertidal estuary located on the border between England 116 117 and Wales in the Eastern Irish Sea. The estuary is tidally-dominated with a mean spring tidal range of 7–8 m (Moore et al., 2009). The geomorphology of the flats causes a tidal asymmetry that is flood 118 119 dominated, resulting in significant accretion of fine sediments in the upper estuary (Halcrow, 2013). 120 Three sites were selected between Hilbre Island and West Kirby (Fig 1), based on the geomorphology 121 of the bed surface at the sampling time. The first site (sandy 1) was dominated by non-cohesive sediment 122 (sand) with wave-influenced 2D current ripples. Site 2 (sandy 2) was similar but had active 2D and 3D 123 ripples. The third site (muddy) was composed of muddy sand, with either a flat bed or relict current 124 ripples (Lichtman et al., 2018).

125 Surface sediment samples were collected at four time points during tidal exposure over three days at each site from 23rd - 31st May 2013 (sandy 1; 23rd - 25th, sandy 2; 26th, 28th - 29th and muddy; 28th - 29th 126 127 and 31<sup>st</sup> May). A full description of the physical conditions during the campaign at the adjacent sites 128 can be found in Lichtman et al., (2018) and the supplementary material. In brief, sampling dates at sandy 1 coincided with the tides transitioning from neaps into peak springs and there was also increased 129 wave action due to high winds on 23<sup>rd</sup>-24<sup>th</sup> May. Despite the strong wave action at sandy 1, which 130 131 caused the maximum wave-current bed shear stresses during wave cycles to be larger, the peak current bed shear stresses during inundation were greater at sandy 2 than sandy 1. Slightly weaker currents 132 133 were observed at the muddy site one day 3, as the tides moved from peak springs toward neaps.

134 Nonetheless, comparable maximum depth-averaged flood/ebb currents were measured across all sites

135  $(0.4-0.7 \text{ m s}^{-1})$  and the maximum water depth at each site ranged between 2 and 3.3 m (Table S1).

136 At each site, 20 surface sediment samples were collected each day ( $n = 5 \times 4$  time points). The first 137 sampling occurred 30-60 min after sediment exposure each day with sampling repeated quarterly during 138 low tide until 60 min before inundation. Samples were collected within 5 m of rigs deployed by NOC 139 Liverpool (Lichtman et al., 2018) and University of Plymouth across an area of approx. 3 m<sup>2</sup>. Sediment 140 cores (2 mm depth, surface area =  $250 \text{ mm}^2$ ) were frozen and stored in liquid nitrogen using the contact 141 core method described in Ford and Honeywill (2002) and Brockmann et al. (2004). Cores were subsequently stored frozen (-80°C) in the dark until processed. To capture both epipelic and 142 143 epipsammic microalgal cells, replicate surface scrapes ( $n = 5, 10 \times 10 \times 2$  mm depth) were collected and stored in 2.5 % w/w glutaraldehyde/filtered seawater solution from time point 2 (T2) only. 144

145 Sample processing – Water content (%) was calculated from wet and freeze-dried core weights before 146 sediment organic matter (SOM, %) was determined by loss-on-ignition at 450°C for 4 h. Chlorophyll a 147 pigments were extracted with 90% acetone following the trichromatic method of Jeffrey & Humphrey 148 (1975). The colloidal and total carbohydrate fractions of the EPS were determined using the phenol-149 sulfuric acid assay (DuBois et al., 1956) following Underwood and Paterson (2003). Due to differences 150 in contents versus concentrations caused by the varied water content of sediment samples, both 151 chlorophyll and carbohydrate measurements are expressed as concentrations per unit area (mg m<sup>-2</sup>, 152 Tolhurst et al., 2005b). The effective particle size distribution (PSD, Grabowski et al., 2012) was 153 determined using a Malvern Mastersizer 2000 laser diffraction analyser (Malvern Instruments Ltd, 154 2013) and summarised using GRADISTAT software (Blott and Pye, 2001) prior to statistical analysis with  $D_{50}$  and mud content (% <63µm) used for further analysis. The relative difference in erosion 155 156 threshold required to suspend a user-defined erosion threshold of 0.01 kg m<sup>-3</sup> was measured using the portable in situ Cohesive Strength Meter (CSM, Paterson, 1989; Tolhurst et al., 1999). In addition to 157 158 the surface erosion threshold, the undrained shear strength was measured using a 33 mm Pilcon shear 159 vane (5 cm depth).

Microphytobenthic community composition – Microphytobenthic cells were extracted from sediment
 scrapes by adopting a modified isopycnic separation technique using silica sol Ludox TM-40®(Ribeiro

et al., 2013). Diatom identification, by means of light microscopy (Zeiss Universal light microscope,
phase and a Ph3-NEOFLUAR oil immersion objective x100 coupled to a 1.0 and 2.0 optivar) followed
Hendey (1964), Hartley et al. (1996) and Round et al. (1990). Cells were identified to species level
where possible and grouped into three ecological guilds (Passy, 2007): (i) "motile" (fast moving and
larger); (ii) "low profile" (short stature, prostrate, adnate erect and slow-moving species), and: (iii)
"high profile" (tall stature, erect, filamentous branched or chain-forming and colonial centrics, largely
non-motile or motile within tubes).

Low Temperature Scanning Electron Microscopy (LTSEM) – Fragments of contact core samples, frozen
in liquid nitrogen (–196.8 °C, 1 atm), were mounted on mechanical stubs and examined using a JOEL
35CF SEM fitted with a LTSEM (Oxford Instruments CT 1500B) following the procedure given in
Paterson (1995).

Statistical analysis - Analyses were performed using "R" statistical software, version 3.1.1 (R 173 Development Core Team, 2014) through the R studio graphical interface (v. 0.98.1083) and in 174 PRIMER software (V.6, PRIMER-E, Ivybridge, UK). Differences in bio-physical variables were 175 176 determined between sites, emersion times and days, after assumptions were tested (Pinheiro et al., 177 2012; Zuur et al., 2007). Necessary transformations were applied to conform to assumptions for 178 parametric statistical testing (ANOVA), where possible or non-parametric Kruskal Wallace (H) tests 179 were employed. Corresponding post-hoc Tukey's or Dunn-sidak tests were applied to detect 180 differences between specific groups. The relationships between the different sediment properties and 181 stability measurements were also assessed using Spearman rank correlations. No significant 182 differences were observed across the different timepoints during emersion for stability or biochemical 183 properties, therefore timepoints were pooled for each day resulting in 20 replicate samples from each 184 day and site. Multivariate analysis of the data using canonical analysis of principal components (CAP) 185 was employed on square-root transformed data to assess the response of multiple bio-physical 186 variables across the sites using constrained ordination taking account of the correlation structure of 187 data (Andersen and Willis, 2003).

188 The MPB community composition between the sites and days was also examined using CAP, based

on Bray-Curtis dissimilarity matrices (Somerfield, 2008). Differences in species between sites and
days were tested using permutational multivariate analysis of variance (PERMANOVA) in addition
to exploring Shannon's diversity index (H') and Pielou's evenness index (Magurran, 2004).

To determine whether the variation in sediment erosion thresholds could be explained by differences in the measured bio-physical properties of the sediment across all sites, data was pooled and distance based linear models (DistLM) were employed (Anderson et al., 2008). Temporal factors (time since emersion and sampling day) were included as explanatory variables along with the various biophysical properties. Marginal and sequential tests were examined using Akaike's information criterion (AICc) and a backwards elimination process to identify the best combination of predictors, that maximised the explained variation with the most parsimonious model (Clarke and Gorley, 2006).

- 199
- 200 Results

Sediment bed properties and stability – The percentage of mud (< 63µm) was significantly higher at 201 the third site (Muddy site, Table 1; 27%,  $H_{(2)} = 93.93$ , P < 0.001) compared to the sandy sites but there 202 203 was no significant difference in mud content between the two sandy sites (0.8 & 1%). For the sandy 204 sites, clean particles were visually observed with very little associated organic matter (Fig 2 A-B) and 205 the grain size distributions were similar (Supp. Fig S1). In contrast, the muddy site exhibited more 206 varied and organic-rich sediments (Fig 2 C-D), but total organic content was relatively low across all 207 sites (<2%) and no significant site differences were detected. The water content, colloidal carbohydrate 208 and chlorophyll a concentrations were all significantly higher at the muddy site compared to both sandy 209 sites (all P < 0.001, Table 1). Differences were between all three sites, with higher contents and 210 concentrations at the muddy site, followed by sandy site 2 and then sandy site 1. The shear strength of 211 the bed was also significantly different between sites, yet this was due to lower strength at sandy site 2, as strengths were similar between sandy site 1 and the muddy site (P > 0.05). 212

At sandy 1, both the colloidal carbohydrates (EPS) and erosion thresholds varied significantly over the sampling days ( $H_{(2)} = 27.12$ , P < 0.001 and  $H_{(2)} = 13.76$ , P < 0.001) but with opposing trends (Table 1). Colloidal carbohydrates were lowest on day two when the erosion threshold was highest, with a decrease in threshold coinciding with an increase in mud and organic content. Overall, the erosion 217 measurements at Sandy 1 correlated very poorly with all measured biological and physical variables, 218 but there was a negative relationship between chlorophyll a concentration (MPB biomass) and the 219 erosion threshold at this site over the sampling days. At the second site (sandy 2), the erosion threshold 220 decreased over the sampling days to its lowest on day 3 (1.8 kPa). This was alongside significant 221 reductions in colloidal carbohydrate concentrations from  $320 \pm 115$  mg m<sup>-2</sup> (day 1) to  $229 \pm 54$  mg m<sup>-2</sup> (day 3, Figure 3;  $F_{2,59} = 6.57$ , P < 0.001). At the muddy site, the erosion threshold was at its highest on 222 223 day one (14.5 kPa) and the lowest on day two (2.7 kPa) also coinciding with the lowest colloidal carbohydrate concentration ( $428 \pm 110 \text{ mg m}^{-2}$ , P < 0.001) and the strongest wave action (Lichtman et 224 225 al., 2018). The  $D_{50}$  of the bed varied very little between days at sandy 1 (Table 1), although there was a 226 small but statistically significant difference at sandy 2,  $D_{50}$  increasing from  $202 \pm 3 \mu m$  on day one to 227  $213 \pm 10 \ \mu\text{m}$  on day three (H<sub>(2)</sub> = 25.88, P < 0.001). No significant differences in D<sub>50</sub> were detected at 228 the sites during different emersion points measured (data not presented).

229 Microalgae community analysis - Significant spatial and temporal differences were detected in the 230 microalgae community across sites, as well as differences between individual days within the sites, (P 231 < 0.01). Diversity (H' index) and evenness (Pielou's index) were significantly different between sites  $(F_{2,24} = 4.91, P < 0.05)$  but differences were relatively small (H' at sandy 1 = 3.14, sandy 2 = 3.32 and 232 muddy site = 3.62, and Pielou's index sandy 1 = 0.92, sandy 2 = 0.87 compared to muddy site = 0.94). 233 234 A few cosmopolitan species such as *Nitzschia frustulum var. inconspicua* were present in almost all 235 samples, across all sites, with Achnanthes punctulata present in greater numbers at the sandy sites. 236 There was a greater abundance of low-profile than motile species at the sandy sites and there were no 237 high-profile species noted. Interestingly, the muddy site had a similar average abundance of low-profile 238 epipelic species to that of the sandy sites, had fewer small epipsammic cells, but had greater numbers 239 of motile and high-profile species that dominated this site. The high proportions of Navicula gregaria 240 (4%), Amphora coffeaeformis var perpusilla (4%) and Pleurosigma aestuarii (4%) at the muddy site 241 appeared to have the greatest effect on site differences in community composition (Figure 4), whereas other smaller species such as Opephora mutabilis (4%) and Cocconeis sp 1 (4%) were more abundant 242 at the sandy sites. Diversity and evenness did not change over time at the sites, but the abundance and 243 244 turnover of key species did vary. At sandy 2, a decline in Opephora mutabilis contributed 9.3% to the overall dissimilarity (42%) between days while the decline in the large motile *Navicula digitoradiata*and *Pleurosigma aestuarii* (6% & 5%, respectively) were the greatest contributors to the overall
dissimilarity (46%) between days at the muddy site.

248 *Bio-physical influences on erodibility* - Inspection of the CAP plots (Figures 5 & 6) and the resulting 249 trace statistic (P < 0.001) confirmed a strong overall difference between the sites based on bio-physical properties. The first axis of the plot (CAP1) was partitioned between several biological variables 250 251 relatively evenly, including chlorophyll a, both carbohydrate fractions, the diversity and number of MPB species present, and water content. Together these variables and the  $D_{50}$  of the bed, which 252 253 exhibited a strong anti-correlation to the biological properties, dominate this axis. On the second axis 254 (CAP2) the  $D_{50}$  of the bed and the undrained shear strength were important factors. A clear spatial 255 separation was observed from the superimposed scatter plot (Figure 5) as well as a temporal component 256 based on draining throughout the emersion period. The close relationships between several variables in 257 the CAP plots were in agreement with correlation analysis (Table 2). As there were no clear dominant 258 biological or physical factors, the majority of properties within both axes were retained for further 259 exploration. Water content was highly correlated to several variables (Table 2) and there were no 260 significant effects of dewatering detected from within each tidal exposure period. Therefore water 261 content was removed from further models. Various single and sequential predictor variables 262 significantly explained the variation in the sediment erodibility across all sites in DistLM (Table 3). 263 When properties were considered individually, both chlorophyll a and organic content significantly 264 explained the greatest variation (at 9% and 8% respectively). While chlorophyll a exhibited a negative 265 effect across all sites, this was primarily driven by the negative relationship at sandy 1, and although 266 MPB diversity and abundance were not good single predictors, they were valuable in sequential tests after consideration of the MPB biomass estimates (chlorophyll a concentration). While the  $D_{50}$  was 267 marginally insignificant in both marginal and sequential tests it was important to retain in the latter yet 268 269 surprisingly the mud content of the sediment was not selected as a good predictor of erosion threshold 270 across the sites.

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272 Discussion

Our results illustrate that the MPB community maintain a key role in sediment dynamics, by surviving 273 274 harsh environmental conditions, and quickly re-establishing biostabilisation. MPB continued to exert an influence on a key ecosystem function; sediment stability across different sediments. Mariotti and 275 276 Fagherazzi (2012) proposed that, given equal intensities of disturbance, the biomass of a biofilm 277 determines whether or not it will be eroded and our results support this. Importantly, our results suggest 278 that this biostabilisation can exert influence on different sediment types. In energetic environments, the 279 frequent turnover and reworking of the sediment may be expected to remove biofilms, hence these 280 habitats are often depicted as abiotic (Figure 7). In very energetic systems, like our sandy site 1, the 281 formation of a fluffy biofilm or layer of cells and EPS may not have create a stable matrix and therefore 282 leads to a 'low biostabilisation' scenario. The lack of incorporation into the bed, explains the negative 283 relationships between key biofilm properties and sediment erosion measurements observed and sandy 284 site 1. However, as grain size was reduced and mud content increased, even slightly, this positively 285 influenced sediment stability, promoting stronger relationships between the biochemical properties of 286 the biofilm and sediment stability at sandy 2 and the muddy site.

287 While frequent resuspension of MPB cells and related EPS may prevent the formation of substantial 288 biofilms, and therefore limit their biostabilisation potential (Aspden et al., 2004), an inoculum often 289 remains in place (Chen et al., 2019) and this can still exert biostabilising effects on the sediment as we 290 have illustrated. This persistence of the biofilm and its stabilising properties means that a biofilm can 291 develop rapidly, if conditions become favourable (see Figure 7; Chen et al., 2019). Previous studies of 292 relationships between EPS carbohydrates and sediment stability have estimated that 2-3 days are 293 required (Lundkvist et al., 2007). However, as we have illustrated in situ that growth does not begin 294 anew at the start of each tidal cycle and biofilms present across different sediment habitats, although 295 invisible, maintain their biostabilisation potential. The stabilising effects of MPB may therefore take 296 less time to develop and become more significant in the natural environment (Chen et al., 2017).

At the sandier sites, the fine sediment and organic matter, which was captured in the suspended sediment traps (data not presented) may have settled onto the sediment surface during slack water, but did not accumulate uniformly on the bed. At sandy 2, a fine organic coating was observed on larger sand grains 300 on ripple crests (Figure 2a) whereas grains from ripple troughs were almost free of OM (Figure 2b) 301 suggesting that OM was being 'caught' from the suspended sediment in the flow. In contrast, the surface 302 sediment at the muddy site was characterised by a much thicker blanket of fine cohesive particles, rich 303 in organic matter and MPB cells, confirmed by LSTEM images (Figure 2 C-D) and biochemical 304 analysis (Table 1). The presence of this organic-rich material can result in positive feedbacks to the 305 system, whereby the adhesive organic EPS and MPB cells trap and bind more fine material, maintaining 306 a muddy bed and reducing suspended sediment concentrations (SSC). This stabilising effect can often 307 be limited to warmer seasons when MPB growth is higher (Borsje et al., 2008) and periods of lower 308 physical disturbance (Widdows and Brinsley, 2002) in temperate regions, but appears to prevail under 309 higher shear stress in this instance. An increase in the  $D_{50}$  of the bed and a decrease in the organic content at the muddy site,  $28^{\text{th}} - 31^{\text{st}}$  was accompanied by a sharp decrease in bed stability (see table 1), 310 311 suggesting the removal of organic material and resulting increase in particle size can destabilise these 312 beds. Organic material creates cohesion between sediment particles, stabilising the sediment when it is 313 bound to particles (Black et al., 2002; Manning et al., 2010; Zhang et al., 2018). However, MPB and 314 OM transported to a particular area during the tide can also form a 'fluff' layer on the bed surface that 315 is easily resuspended if it is not incorporated into a biofilm (Orvain et al., 2003). This is likely the case 316 at sandy site 1 as high MPB biomass (chlorophyll a concentration) and EPS were observed on days 317 when the erosion threshold was reduced. The cells and EPS detected in the sediment surface were 318 therefore unlikely to have formed a protective film. Indeed, in sandy sediments (like sandy site 1), small 319 diatoms tend to attach themselves to the grains and coat individual sand grains in EPS rather than 320 forming a substantial biofilm per se.

The different mechanisms by which MPB and EPS develop in dynamic sandy sites may explain the negative relationship between chlorophyll a and erosion threshold at sandy site 1. Substantially more EPS is produced and excreted by epipelic (motile) diatoms, like the taxa dominating the muddy site. While previously it has been thought that >50% of the MPB community must be epipelic species (Underwood et al., 1995; Underwood & Paterson, 2003), in the Dee Estuary it appears that the proportion may be much lower. These differences in the relationships often hinder attempts to generalise MPB biomass effects on erodibility, and lead to significant differences between studies. 328 Despite the importance of accurately forecasting erosion threshold parameters for sediment transport 329 predictions (Sanford, 2008) the influence of biological cohesion across different habitats is rarely 330 considered in these models (Le Hir et al., 2007). This is primarily due to the complexity of intertidal 331 systems and differences in biological and physical processes across sediment gradients. There is likely 332 a threshold of development under which very coarse sandy sites may not be positively influenced by MPB biostabilisation, or under extreme conditions like the significant wave action at sandy site 1 prior 333 334 to sampling (Lichtman et al., 2018). Nonetheless, other sandy sites (such as our sandy site 2) can be 335 positively influenced by biostabilisation (Larson et al., 2009). Cells and EPS that are not incorporated 336 as a biofilm can be easily suspended and recorded as erosion by the CSM system and these results 337 suggest that characterising the MPB community can help to explain these differences (Figure 7A).

338 In this study, EPS (as colloidal carbohydrates) were positively related to sediment erosion thresholds, 339 even at sandy site 2. At this site, the community was composed primarily of low profile pioneer species, 340 and had a limited number of motile forms (discussed further in microalgal communities section). This 341 relationship suggests that even when the EPS matrix does not form a substantial biofilm on the surface, 342 it still offers some form of protection to the underlying sediment (Figure 7B). Laboratory (Malarkey et 343 al., 2015; Parsons et al., 2016) and field investigations (Baas et al., 2019; Lichtman et al., 2018) have 344 recently illustrated the influence of low EPS contents distributed deeper into the sediment bed. In these 345 investigations microbially-produced EPS hampered sediment transport, bedform development, and 346 bedform migration without the presence of a visible biofilm on the surface (see also Chen et al., 2017). 347 These vectors of change are undoubtedly part of the short-term variation due to changes in the 348 spring/neap cycle and daily weather fluctuations, but these changes can also be the first steps toward a 349 transition towards an alternative state (Van de Koppel et al., 2001).

350

351 *The microalgal communities* 

352 Over the relatively short sampling period described here, the microalgae community at the sandy sites 353 appeared to remain in early successional stages, whereas the community at the muddy site had already 354 developed into a more vertically-structured community composed of stalked, filamentous and motile 355 microalgae (Winsborough and Golubic, 1987). The latter forms can withstand stronger flow velocities 356 and abrasion from moving sediments by migrating into the muddy sediment or creating filaments. 357 However, adnate forms such as Achnanthidium are well equipped to resist flow (Passy, 2007). Certain 358 diatom species are indicators of the flow regime, with particular species exhibiting preference for high 359 flow such as Achnanthidium spp (Passy, 2007). While flow differences can result in different 360 communities, the effects can also be dampened by differences in turbidity (Soininen, 2004). 361 Nonetheless, information on community composition together with information on the bio-physical 362 properties of the sediment can be useful for determining the differences in erosional resistance and 363 potential biostabilisation across different habitats. Our results indicate the number of species was 364 altered over the days at the different sites and species composition changed over short temporal scales. 365 However, in each site the different guilds (low profile, high profile, motile etc) remained dominant due 366 to their adaptations to the flow environment. While MPB community succession can be revealed 367 through the microscopic identification of cells, this is time-consuming and could be complimented by 368 next generation sequencing of the prokaryote and eukaryote communities for longer-term studies (Hicks 369 et al., 2018). This would provide a more comprehensive microbial community analysis in relation to 370 biostabilisation (Paterson et al., 2018) as the diversity of prokaryotes has been linked to hydrodynamic 371 regimes (Besemer et al., 2009). Such an approach would be incredibly useful for capturing the 372 transformation of sites that are frequently disturbed and dynamic in nature, into more stable muddy 373 habitats over longer timescales. MPB composition and structure can reflect differences in flow regimes 374 (Krajenbrink et al., 2019) and as the community changes, primary productivity and the production of 375 EPS exudates will vary, with knock on effects on various ecosystem structure and functions (Hope et 376 al., 2019). As the hydrodynamic effects on MPB communities can modulate the effects of others 377 stressors (Polst et al., 2018; Villeneuve et al., 2011) understanding the interaction between the 378 community and hydrodynamics across different sediment habitats is essential. For instance, Achnanthes 379 spp and Nitzschia inconspicua, were observed in all Dee Estuary sites. These cells are often one of the 380 first species to inhabit recently disturbed sediment (Cardinale, 2011), and are cosmopolitan (Sabater, 381 2000). They can grow prostrate to the surface or adnately (Berthon et al., 2011; Cardinale, 2011), which helps them withstand high flow velocities (Passy, 2007). We have however illustrated that these are not 382 383 displaced by the development of the biofilm as they were still present in our muddier site. These could

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be key species to examine for the effects of other stressors in these environments. These pioneers promote the rapid recolonization of the sediment bed after physical disturbance, instigating a biological succession, which promotes greater biodiversity and heterogeneity in the biofilm and among higher organisms (Balvanera et al., 2006). This can pave the way for a more heterogeneous community of microorganisms and a complex habitat that can increase biostabilisation (Paterson et al., 2018).

The survival of algal cells during tidal inundation or their deposition from the water column establishes 389 390 a potential for microbial growth and eventual biofilm formation if conditions allow (Figure 7B-C). This 391 was evident from the differences between sites in the Dee Estuary. At sandy site 1, the MPB community 392 was composed of pioneer species that turned over with the prevailing hydrodynamic conditions. At the 393 opposite end of the spectrum, the MPB community at the muddy site was more stable and composed of 394 larger epipelic species. However, at sandy site 2, the community was distinct and appeared to be 395 intermediate between the two other sites. In these transitioning sites, the maximum variation in the 396 sediment surface erosion is expected (Figure 7B) as the surface is patchy. As MPB communities develop 397 and grow on the sediment surface, this drives the capture of more cohesive material and this positive 398 feedback enhances the development of a more homogenous and stable surface dominated by biofilms 399 (Figure 7C).

400 Disturbance from tidal flow can exert the same effects as large bioturbating fauna oxygenating the 401 sediment surface layers (Huettel et al., 2003; Precht and Huettel, 2003). These processes are important 402 for soft sediment ecosystem functions such as sediment oxygenation, biogeochemical cycling and, 403 depending on the organic enrichment of the sediment, degradation processes (Widdicombe and Austen, 404 2001). The contribution of large infauna has recently be discussed elsewhere (Hillman et al., 2019) and 405 are of course important to consider in many habitats. Low numbers of large fauna were observed at 406 these sites, therefore we focussed on the physical processes and the interaction with microbial organisms that are known to stabilise and disturb the bed. The close spatial association of visibly 407 408 different sedimentological properties suggests bio-physical factors may contribute to the variation over 409 short distances despite the similarity of dynamic context.

Understanding the bio-physical factors influencing sediment stability across different habitats allowsus to begin to discern how and why mixed beds occur and the mechanisms by which they alternate

412 between muddy, mixed and sandy habitats. Muddy sites can capture fine sediment, due to the cohesive 413 nature of fine particles, MPB and EPS in the surface layers (Table 1). This cohesiveness can prevent 414 fine particles from winnowing during inundation and result in higher erosion thresholds. These bio-415 physical properties can lead to the formation of a cohesive matrix that can effectively trap additional 416 material from the water column and improve the clarity of the overlying water. Positive correlations between mud content, organic content, EPS carbohydrates, and MPB biomass/community indices were 417 418 apparent at the muddy site (Figure 5) indicative of biofilm development where higher numbers of motile 419 diatoms were present. It has previously been suggested that relatively high proportions of motile diatoms, and hence high EPS concentrations, are required to trap new deposits of sediment (Underwood, 420 421 1997) and counteract the physical forces that resuspend sediment, and this can lead to positive 422 feedbacks. Van de Koppel et al. (2001) highlighted these feedbacks, and proposed that ecosystem 423 engineering (Jones et al., 1994), principally by MPB, can mediate changes in bed sedimentology.

424 At the sandy habitats, the regular physical disturbance from waves and currents over the tidal cycle 425 prevented the accumulation of larger MPB, which limits biofilm development (Blanchard et al., 2001). Over the course of this relatively high-resolution investigation, the  $D_{50}$  of the sandy site 1 increased 426 427 despite the fine nature of the material frequently collected in suspension traps at the sites (data not 428 presented). This was primarily due to the prevailing wind wave action at this site during this period 429 (Lichtman et al., 2018; Table S1). Fine cohesive sediment has to be removed frequently in sandier 430 habitats, through resuspension or winnowing to impede the development of biofilm growth but MPB 431 are still present and still exert influence over the sediment dynamics. When conditions are altered this 432 can allow the MPB to proliferate, significantly increasing the erosion threshold and instigating a 433 transition to finer sediment. These differences in the erosive nature of the bed and the fate of settling 434 material, is key to maintaining the differentiation between patches and increases overall habitat heterogeneity (Weerman et al., 2011) and functioning of soft sediment ecosystems (Thrush et al., 2008). 435 436 By investigating the short term, temporal dynamics influencing the MPB community, and the feedbacks 437 between the biomass, community composition, exudates of MPB and biostabilisation potential we can 438 begin to understand the conditions required to instigate the changes that lead to transitions and postulate

how the microbial organisms in these habitats can persist and continue to exert an influence on sedimentstability.

441

442 Conclusion

443 The relative influence of MPB and EPS on sediment stability and transport remain poorly understood across different sediment habitats. The results of the study suggest various biological properties of the 444 445 bed associated with the MPB significantly influence the short-term variability in the erodibility of different surface sediments. Importantly, we illustrate that while MPB diversity explained very little 446 447 variation in marginal distance based linear tests, primary producer diversity was a significant predictor 448 when MPB biomass was also considered in sequential tests. We emphasise the importance of 449 considering the microbial diversity when assessing their influence on ecosystem functions such as 450 sediment stability. Further evidence of biological cohesion across natural habitats of increasing 451 complexity and at multiple spatial and temporal scales is required in order to understand the biological influence on sediment dynamics. Further data with natural gradients of sand and mud should be 452 453 examined and the influence of larger benthic organisms included to document the influence of 454 biological properties across different habitats, under differing physical conditions and with increasingly 455 complex communities. This will facilitate the use of these variables in future sediment transport models.

456

457 *Acknowledgements* 

458 This work was supported by the UK Natural Environment Research Council (NERC), grant code 459 NE/I027223/1 (COHBED). DMP received funding from the Marine Alliance for Science and 460 Technology for Scotland (MASTS), funded by the Scottish Funding Council (grant reference 461 HR09011). JM and DRP were partially funded by a European Research Council Consolidator Award (725955). We are grateful to the NOC Ocean Technology and Engineering group, and field technicians 462 463 from the University of Plymouth, Bangor University and St Andrews for instrument set up and deployment and sample collection. Thank you to Emily Carr and Stefano Schenone for their help with 464 graphics. LTSEMs were produced by Mr Irvine Davidson, University of St Andrews, as his last SEM 465 466 work before retirement. His phenomenal contributions over the years are greatly appreciated. We also

467 wish to thank the editor, associated editor and two anonymous reviewers for their valuable comments

468 and suggestions on an earlier version of this manuscript. The authors have no conflicts of interest

- regarding this work.
- 470
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	Colloidal	Chl a	Water	Organic	$D_{50}$ Bed	Mud content	Sed. erosion	Shear Strengtl
	carb conc (mg m <sup>-2</sup> )	conc (mg m <sup>-2</sup> )	content (%)	content (%)	(µm)	Bed (%)	threshold (kPa)	(kPa)
Sandy site 1	159	13	18	1.5	223	0.8	12.1	14.5
	± 87	± 3	± 1	± 1.1	± 6	± 1.8	(8.8 – 17.5)	(8.9 – 16.7)
23 <sup>rd</sup> (D1)	$\begin{array}{c} 235 \\ \pm  41 \end{array}$	16 ±2	18 ± 1	0.8 ± 0.3	222 ± 5	0.0 ± 0.0	11.0 (8.8 – 14.8)	15.0 (14.1 - 16.0)
24 <sup>th</sup> (D2)	88	10	18	2.8	221	1.7	18.6	14.0
	± 67	±2	± 1	± 2.8	± 7	± 2.2	(13.2 - 25.2)	(12.0 - 15.5)
25 <sup>th</sup> (D3)	147	13	18	0.8	225	1.3	11.0	14.5
	± 74	±2	± 1	± 0.2	± 3	± 1.9	(6.6 – 17.0)	(8.9 - 16.7)
Sandy site 2	269	26	19.9	0.8	204	1.0	4.4	6.8
	± 88	±4	± 0.9	± 0.2	± 9	± 2.0	(1.8 - 6.7)	(5.5 – 12.5)
26 <sup>th</sup> (D1)	320	28	20.4	1.0	202	2.1	6.6	6.2
	± 115	± 3	± 0.8	± 0.9	± 3	± 2.6	(2.2 – 14.8)	(6.0 - 7.0)
28 <sup>th</sup> (D2)	255	23	20.1	0.7	197	0.0	4.9	6.5
	± 59	±4	± 1.2	± 0.2	± 4	± 0.0	(2.2 - 9.0)	(6.1 – 7.9)
29 <sup>th</sup> (D3)	229	26	19.6	0.9	213	1.9	1.8	10.2
	± 54	± 3	± 0.5	± 0.1	± 10	± 2.4	(1.8 – 1.8)	(7.1 – 11.4)
Muddy site	532	44	21.0	1.4	156	27.0	5.4	13.2
	± 165	± 11	± 1.3	± 0.5	± 23	± 7.0	(3.3 – 11.5)	(9.0 – 18.0)
28 <sup>th</sup> (D1)	609	44	21.6	1.5	140	30.8	14.5	14.0
	± 186	± 14	± 1.2	± 0.4	± 19	± 5.0	(7.7 – 23.4)	(10.5 – 1.2)
29 <sup>th</sup> (D2)	428 ± 110	40 ± 8	21.1 ± 1.5	1.3 ± 0.4	159 ± 15	27.1 ± 3.9	2.7 (1.8 - 5.0)	
31 <sup>st</sup> (D3)	557	47	21.0	1.6	169	22.8	5.5	13.0
	± 141	±9	± 1.3	± 0.6	± 24	± 8.2	(2.7 – 8.8)	(11.7 - 13.5)

Table 1: Temporal variation in the mean / median values of biological and physical measurements at the sandy site and muddy site for all days combined and then each individual day of sampling. Top number = mean / median value, bottom number = Standard deviation / interquartile range.

# Limnology and Oceanography

Table 2: Spearman rank correlation coefficients for all variables within and across sites. 1 <sup>st</sup> number = sandy 1, 2 <sup>nd</sup> = sandy 2, 3 <sup>rd</sup> = muddy site, bottom number
( <b>bold</b> ) = all sites. Sig levels - '***' = $P < 0.001$ , '**' = $P < 0.01$ , '*' = $P < 0.05$ = no significant correlation detected. $N = 60$ per site.

	Colloid carbs (mg m <sup>-2</sup> )	Chl a (mg m <sup>-2</sup> )	Water content (%)	Organic content (%)	D <sub>50</sub> (μm)	Mud content (%)	Erosion threshold (kPa)	Undrained shear strength (kPa)
Colloid carbs (mg m <sup>-2</sup> )	-							
Chl a (mg m <sup>-2</sup> )	0.62*** 0.64*** 0.57*** <b>0.83</b> ***	-						
Water content (%)	- 0.44*** <b>0.52</b> ***	0.27* 0.33** 0.59*** <b>0.62</b> ***	-					
Organic content (%)	- 0.34** <b>0.65</b> ***	- 0.40** <b>0.68</b> ***	-0.36** - 0.43*** <b>0.42</b> ***	-				
D <sub>50</sub> (μm)	- - - <b>0.70</b> ***	- - 0.62***	- - -0.26* - <b>0.50</b> ***	- 0.39** - - <b>0.55</b> ***	-			
Mud content (%)	- - 0.72***	- - - 0.69***	- 0.25* <b>0.47</b> ***	- - 0.63***	0.42*** 0.44*** -0.94*** - <b>0.85</b> ***	-		
Erosion threshold (kPa)	- 0.35** 0.28* <b>0.35</b> ***	-0.43*** - - 0.18*	- - 0.20*	-0.31*	- 0.26* <b>0.29</b> **	- - 0.21*	-	
Undrained shear strength kPa)	- - 0.65***	- - 0.57***	-0.3* -0.61*** - <b>0.20</b> *	- 0.30* - <b>0.68</b> ***	- 0.36** - - <b>0.64</b> ***	- - 0.72***	-0.4** - -	-

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Table 3: The % variation in the erosion threshold of the sediments across all sites, explained by various bio-physical properties. Both marginal (single predictor) and step-wise sequential results for DistLM are presented. Significance levels indicated as '\*\*\*' = P < 0.001, '\*\*' = P < 0.01, '\*' = P < 0.05, <sup>¥</sup> = marginally insignificant P<0.10 and NS = not significant.

	AICc	Pseudo-F	Expl. Variation (%)	Cumul. Expl. Variation (%)
Marginal tests				
Chl a (mg m <sup>2</sup> )		17.32	9***	
Shannon (H) index of MPB diversity		0.34	2***	
MPB species abundance		0.77	4 <sup>NS</sup>	
Organic content (%)		14.81	8***	
Colloidal carbs (mg m <sup>2</sup> )		3.25	2 <sup>¥</sup>	
<i>D</i> <sub>50</sub> (μm)		3.58	2 <sup>¥</sup>	
Sequential tests				
Chl a (mg m <sup>2</sup> )	-13.63	17.32	9***	9
Shannon (H) index of MPB diversity	-32.20	21.53	10***	19
MPB species abundance	-42.92	13.00	6***	25
Organic content (%)	-50.85	10.05	4**	29
Colloidal carbs (mg m <sup>2</sup> )	-56.94	8.14	3**	32
<i>D</i> <sub>50</sub> (μm)	-58.19	3.32	2¥	34



Figure 1: The location of the sampling sites on the intertidal flats near West Kirby and Hilbre Island. Inset -Position of the Dee Estuary, near West Kirby Liverpool, England.

159x89mm (96 x 96 DPI)



Figure 2: Low temperature scanning electron micrographs (LTSEM) of the intact sediment surface at A) Sandy site 2, crest of ripples B) Sandy site 2, troughs C) Muddy site, general surface and D) Muddy site, close-up image of organic material between sediment grains. Scale bars:  $100\mu m$  for A-C and  $10\mu m$  for D.

166x166mm (96 x 96 DPI)





135x152mm (96 x 96 DPI)



Figure 4: Low-temperature scanning electron micrographs of diatoms harvested from the surface sediment of the Dee Estuary. Scale bars = 1µm/10µm where stated. A) Opephora mutabilis (Grunow), B) Epipsammic sp 1, C) epipsammic cell embedded in a sediment particle in a matrix of EPS, D) Navicula gregaria (Donkin), E) Planothidium haukiana (Grunow), F) Nitzschia sp 1, G) Amphora coffeaeformis var coffeaeformis (Agardh) Kützing, H) Pleurosigma aestuarii, (Brébisson ex Kützing) and several small epipsammic cells / sediment particles, I) Cocconeis peltoides (Hustedt), J) Amphora tenerrima (Aleem & Hustedt), K) Thalassionema spp (Grunow) L) Diploneis spp.

186x228mm (96 x 96 DPI)



Figure 5: Canonical analysis of principal components (CAP) plot of euclidean distance similarities between samples. The correlation circle overlays measured variables that were influencing the similarity/dissimilatory between the samples. All data was square root transformed and normalised prior to analysis. n = 60. D50 - D50 of the particle size distribution, top\_vane – undrained shear strength, TP – time point since emersion, H' - Shannon diversity index, OC - organic content of sediment (%), MC – Mud content of sediment (%), C/carbs - colloidal carbohydrate concentrations, chl a - chlorophyll a concentrations, WC - water content (%).

159x119mm (96 x 96 DPI)



Figure 6: Canonical analysis of principal components (CAP) plot of Bray Curtis similarities in the microalgae community composition of the sediment surface between the two sandy sites (S1 and S2) and the muddy (M) site over three days at the Dee estuary.

159x105mm (96 x 96 DPI)



Supplementary figure S1: Grain size distributions for A) sandy site 1. B) Sandy site 2. C) Muddy site. Values are averaged across all days at each site. Note that the y-axis of the muddy site is on a different scale.

### Limnology and Oceanography

Table S1: Peak hydrodynamic values during the inundation at the three sites based on the measurements of Lichtman et al. (2018). The current and wave-current bed shear stresses are the mean and maximum during a wave cycle (the closer they are to one another the weaker the effect of the wave).

Date	Site	Water depth	Depth-averaged current	Significant wave	Wave period	current bed shear stress	Wave-current
		(m)	(m s <sup>-1</sup> )	height	(s)	(Nm <sup>-2</sup> )	bed shear stress
				(m)			(Nm <sup>-2</sup> )
23/05	S1	2.23	0.49	0.38	7.8	0.41	1.09
24/05	S1	2.59	0.63	0.48	7.9	0.62	1.38
25/05	S1	2.88	0.56	0.11	8.3	0.51	0.59
26/05	S2	3.09	0.65	0.27	4.8	0.72	0.79
28/05	S2, M	3.34	0.71	0.28	6.1	0.84	0.93
29/05	S2, M	3.01	0.62	0.18	6.4	0.65	0.73
31/05	М	2.08	0.46	0.14	4.6	0.38	0.39

N. I.



Figure 7. Conceptual diagram of microbially-induced variability in surface sediment erodibility. A) Noncohesive sediment lacking biogenic influence that in theory would show little variability in surface behaviour forms a predictable and homogenous habitat. B) Colonisation of the non-cohesive bed by microbial cells producing extracellular polymeric substances (EPS) and the initial growth of small microbial colonies creates heterogeneity in the localised surface response to shear stress. Increasing the local resistance to erosion in some patches. C) A fully colonised substratum where biofilm development has created a more uniform sediment surface, once again reduces the variability of the system but further increases the erosional resistance. Top left: Spatial variation in erosion resistance across the bed. Condition A = homogenous abiotic grains producing constant and predictable erosion thresholds. Condition B = A highly heterogeneous system with an erosion threshold influenced by the complexity of local conditions, and the patchy distribution of MPB and bacterial biofilms. Condition C = coherent biofilm increases sediment stability and reduces erodibility in a consistent manner across the bed. This creates a more homogenous response to erosional stress until bed failure. Erosive loss from areas of biofilm growth (C), can lead to the depositional gain of MPB at other sites (B). This may lead to the development of a substantial biofilm (C), or the subsequent resuspension of the MPB again. At more energetic sites (A), fine sediment and MPB are deposited during slack tide, but these are resuspended on the next tide, maintaining a relatively homogenous system where MPB may be present but there is no stabilising effect of the biofilm due to frequent resuspension. These states may alter as local conditions change including seasonal, light nutrient and temperature differences (which would stimulate the MPB and biofilm growth), and hydrodynamic conditions which increase erosional stress on the surface sediment.

275x190mm (96 x 96 DPI)