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1 Highlights

- 2 • Flood and drought events had negative impact on indicators of ecosystem function
- 3 • This grassland was more resistant and resilient to drought than flood
- 4 • Flooding led to pronounced and persistent shift in plant and microbial communities
- 5 • The combination of flood and drought stress increased the resilience of the system

6

7 **Resilience of ecosystem service delivery in grasslands in response to single and**
8 **compound extreme weather events**

9

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32 **Abstract**

33 Extreme weather events are increasing in frequency and magnitude with profound effects on
34 ecosystem functioning. Further, there is now a greater likelihood that multiple extreme events
35 are occurring within a single year. Here we investigated the effect of a single drought, flood
36 or compound (flood+drought) extreme event on temperate grassland ecosystem processes in a
37 field experiment. To assess system resistance and resilience, we studied changes in a wide
38 range of above- and below-ground indicators (plant diversity and productivity, greenhouse
39 gas emissions, soil chemical, physical and biological metrics) during the 8 week stress events
40 and then for 2 years post-stress. We hypothesized that agricultural grasslands would have
41 different degrees of resistance and resilience to flood and drought stress. We also investigated
42 two alternative hypotheses that the combined flood+drought treatment would either, (A)
43 promote ecosystem resilience through more rapid recovery of soil moisture conditions or (B)
44 exacerbate the impact of the single flood or drought event. Our results showed that flooding
45 had a much greater effect than drought on ecosystem processes and that the grassland was
46 more resistant and resilient to drought than to flood. The immediate impact of flooding on all
47 indicators was negative, especially for those related to production, and climate and water
48 regulation. Flooding stress caused pronounced and persistent shifts in soil microbial and plant
49 communities with large implications for nutrient cycling and long-term ecosystem function.
50 The compound flood+drought treatment failed to show a more severe impact than the single
51 extreme events. Rather, there was an indication of quicker recovery of soil and microbial
52 parameters suggesting greater resilience in line with hypothesis (A). This study clearly
53 reveals that contrasting extreme weather events differentially affect grassland ecosystem
54 function but that concurrent events of a contrasting nature may promote ecosystem resilience
55 to future stress.

56 **Keywords:** Climate change, soil quality, drought, flooding, greenhouse gas emissions,
57 microbial community

58

59 **1. Introduction**

60 Grasslands represent an important global agroecosystem covering an area of ca. 11 million
61 km² and 7 % of the earth's terrestrial area (O'Mara, 2012). In addition to food production,
62 grasslands provide a wide range of ecosystem services (ES), such as biodiversity provision,
63 climate regulation and natural hazard protection (MEA, 2005; O'Mara, 2012) which rely on
64 effective belowground soil functioning (Adhikari and Hartemink, 2016; Brussaard, 1997;
65 Wagg et al., 2014). The effects of climate change are already being felt, with the frequency
66 and severity of extreme weather events increasing in every region of the globe (IPCC, 2021).
67 Such extreme events have been shown to negatively affect grassland yields and biomass
68 production (Environment Agency, 2006; Niu et al., 2014). Recent spatial analysis of long-
69 term temperature and precipitation records, in combination with land use data, highlighted an
70 increasing risk of flooding and drought (single and in combination/succession) in many UK
71 pastoral landscapes (Dodd et al., 2021). To safeguard the vital ES provided by agricultural
72 grasslands, we need to fully understand the impact of such events on the whole ecosystem.

73 Precipitation regimes regulate grassland ecosystem structure and function due to the
74 influence on soil moisture, a major driver of plant growth and microbial activity (Bloor and
75 Bardgett, 2012). While there is some uncertainty in the response pattern (Evans et al., 2022),
76 plants and microbes are considered to exhibit a hump-shaped response to moisture with
77 substrate diffusivity and oxygen limitation constraining growth at the low and high extremes
78 respectively. Consequently, extreme events which rapidly and dramatically change the
79 moisture conditions (such as flood and drought) have the potential to cause a large-scale
80 change in above and below ground ecosystem processes. A wealth of research on drought has

81 demonstrated large scale changes in ecosystem properties including reductions in primary
82 production (Ciais et al., 2005; Hoover et al., 2014), changes in above and below ground
83 community structure (de Vries et al., 2018; Knapp et al., 2020; Ochoa-Hueso et al., 2020;
84 Ochoa-Hueso et al., 2018) and altered C and nutrient cycling (Bloor and Bardgett, 2012;
85 Deng et al., 2021; Dijkstra et al., 2015). However, research into ecosystem responses to
86 prolonged flooding at the field scale have been largely neglected, and while some
87 experiments include increased precipitation regimes, these rarely examine prolonged flooding
88 (Abbasi et al., 2020), despite evidence of widespread impacts on vegetation across the globe
89 (Famiglietti et al., 2021). Flooding disruption of nutrient cycles (Sánchez-Rodríguez et al.,
90 2019a,b) and increased N₂O emissions (Oram et al., 2021; Sánchez-Rodríguez et al., 2019a;
91 Sánchez-Rodríguez et al., 2019b) along with changes in microbial communities (Sánchez-
92 Rodríguez et al., 2019a,b; Unger et al., 2009) have been observed in mesocosm studies under
93 controlled conditions. However, such studies may exhibit differing responses to the field
94 scale (Unger et al., 2009). Alternatively, flood impacts have been assessed following natural
95 flood events on existing field trials (González Macé et al., 2016; Harvey et al., 2019; Wagner
96 et al., 2015). Nevertheless, these opportunistic trials often have limited ability to regulate the
97 experimental conditions and are often poorly instrumented and lack appropriate controls and
98 plot scale flood experiments remain a key gap in the literature.

99 In addition to single extreme events, there is growing recognition that many ecosystems
100 can experience a combination of multiple climatic drivers or hazards often termed compound
101 extremes (Zscheischler et al., 2018). Such compound events are likely to have a larger effect
102 on ecosystem function, through a reduction in the systems resilience, amplifying the impact
103 and potentially leading to bi-modal response-recovery patterns indicative of a regime shift
104 (Rillig et al., 2019). While research into the impact of compound events on ecosystems is
105 growing, these have been mainly focused on combinations of drivers with additive properties,

106 for example heatwaves and droughts (Arain et al., 2022) or storm surge and flooding (Wahl
107 et al., 2015). However, recent research suggests that the occurrence of multiple extreme
108 events, with contrasting drivers, within the same year is also increasing; for example a swift
109 transition from flood to drought events or *vice versa* (Dodd et al., 2021).

110 As ecosystem function is highly dependent on soil moisture, contrasting events of
111 extreme precipitation amount, occurring in quick succession, are likely to have a different
112 impact to each event alone. A deluge event during summer drought was seen to alleviate
113 reductions in plant production and soil respiration in a natural semi-arid grassland (Post and
114 Knapp, 2020) and numerous studies have examined the effect of rewetting of dry soils on many
115 aspects of ecosystem function (Birch, 1958; Borken and Matzner, 2009; Fierer and Schimel,
116 2002; Gordon et al., 2008). However, there has been little investigation of the reverse, where
117 rapid drying of wet soils may occur due to reduced rainfall. The frequency of winter and spring
118 flooding and that of summer droughts is predicted to increase (Fowler and Kilsby, 2003;
119 Thompson et al., 2017). Furthermore, as extreme events are becoming more frequent and
120 intense (Li et al., 2019; Trenberth, 2011), the likelihood of a second extreme occurring before
121 the system has fully recovered also increases, potentially magnifying the impact (Schwalm et
122 al., 2017) and tipping the system to a new functional state. We propose two alternative
123 hypotheses of ecosystem response to this scenario: (A) the drawdown of soil moisture may be
124 slower in previously flooded soils promoting resilience to subsequent drought and a reduction
125 in observed ecosystem impact (Post and Knapp, 2020), or (B) if the drought event occurs prior
126 to ecosystem recovery, a second stress event may lead to adverse impacts on biotic
127 communities and potential community collapse, with multi-year impacts on ecosystem
128 function.

129 Our study directly addresses gaps in knowledge surrounding the short and medium-
130 term impact of simulated severe drought and flood events on the resilience of grassland

131 ecosystems at the field plot scale. Further, we assess the impact of compound weather events
132 characterised by a spring flood followed by a summer drought. Here, we use a set of above
133 and below ground biotic and abiotic indicators to quantify the effect of these extreme events
134 on ecosystem process of six essential grassland ES: forage production, nutrient cycling,
135 organic matter decomposition, climate regulation, pollution regulation and biodiversity
136 provision. Changes in these indicators were used to assess the impact of each event type
137 immediately post stress (reflecting system resistance) and the legacy implications for longer-
138 term ecosystem function through monitoring indicator recovery (reflecting system resilience).
139 We hypothesised that (1) extreme weather events will negatively impact on ecosystem
140 processes; (2) grassland ecosystems will have different degrees of resistance and resilience to
141 flood or drought stress. Additionally, we investigated the two alternative hypotheses
142 presented above (A and B) for the combined flood and drought event.

143

144 **2. Materials and methods**

145 *2.1. Treatments and plot establishment*

146 The study site was located at the Henfaes Agricultural Research Station,
147 Abergwyngregyn, Gwynedd, North Wales, UK (53°14'21"N, 4°00'57"W) on sheep-grazed,
148 *Lolium perenne* L. dominated, low intensity grassland. The soil at the site is a sandy loam
149 textured Eutric Cambisol (Typic Hapludalf) overlying a mixed glacial till parent material.
150 The site has a temperate-oceanic climate with a mean annual rainfall of 1060 mm and
151 temperature of 10 °C.

152 Sixteen 3 m × 3 m plots were established in winter 2015. The experiment consisted of four
153 treatments and was conducted in three phases. The treatments consisted of (1) spring flood
154 (flood), (2) summer drought (drought), (3) spring flood followed by summer drought
155 (flood+drought), and (4) ambient conditions (control). Each treatment had four replicates ($n =$

156 4) arranged in a randomised plot design. The three data collection phases were (1) flood
157 period, (2) drought period and (3) recovery.

158 Although the site was previously grazed with sheep (*Ovis aries* L.), livestock were
159 excluded from the plots during the construction phase. Two weeks prior to the start of the
160 trial (April 2016), sheep were briefly reintroduced (0.22 sheep m⁻²; 2 d) to mimic mob
161 grazing and trampling in this sheep-based pasture system.

162 To prevent lateral water escape, the eight flood plots were hydrologically isolated.
163 This was achieved by vertically inserting 80 cm deep PVC plastic boards (5 cm thick; Tatra-
164 Rotalac Ltd., Manchester, UK) lined with an impermeable butyl rubber membrane
165 (Pondkeepers Pond Liners Ltd., Billingham, UK; Fig. S1) around the plot boundary. The
166 boards were buried to a depth of 30 cm in the soil, leaving 50 cm aboveground to hold the
167 flood water. A physical barrier was also inserted to a similar depth to isolate the eight non-
168 flood plots to ensure a similar level of physical disturbance.

169 To mimic on-farm management practices, lime (3 t ha⁻¹) and NPK compound fertiliser
170 (25:5:5; 50 kg N ha⁻¹) were applied to all plots in April 2016 (14 d prior to flooding), in line
171 with past fertiliser application rates in this field. The same rate of NPK fertiliser was applied
172 again in August 2016 during the drought period.

173

174 2.2. *Experimental details*

175 The flooding treatments were designed to simulate a fluvial flood event, reflecting the
176 scale and duration of similar events which have occurred within the local and wider region in
177 the last decade (Defra, 2014; Harvey et al., 2019; Huntingford et al., 2014; Slingo et al.,
178 2014). However, it should be noted that the site had no previous history of flooding. Figure
179 S1 shows the layout of the field trial.

180 Flood phase: In spring 2016, eight flood plots were submerged with water from the
181 adjacent river to a depth of 20 cm, which was sufficient to fully submerge all vegetation (Fig.
182 S1.c). Sediment was added to each of the plots following the initial flooding to a depth of 2-3
183 mm to simulate the deposition of eroded soil and sediment typically associated with river
184 flooding and observed in recent extreme flood events (ADAS, 2014). This was achieved by
185 suspending 13.5 kg of surface soil in the floodwater. The soil was taken adjacent to the plots
186 (0-7 cm depth) under the same management regime.

187 The plots were maintained at a constant flood level of 20 cm for eight weeks during
188 April to June, by topping up the flood water, via a ball float valve connected to a reservoir of
189 river water, reflecting the unprecedented flooding events observed across UK agroecosystems
190 in the winter of 2013 - 2014 (Defra, 2014; Slingo et al., 2014). At the end of the flood period
191 the floodwater was removed using a pump over a 4 h period.

192 Recovery phase 1: All plots were left under ambient conditions for four weeks
193 following the end of the flood phase and before the initiation of the drought phase.

194 Drought phase: At the end of recovery phase 1, rain-out shelters (4 m × 4 m area)
195 were erected over eight plots, four of which were previously subjected to the spring flood.
196 The rain-out shelters were constructed with PalSun[®] polycarbonate sheets (2 mm thick;
197 Plastock Ltd, High Wycombe, UK), mounted on wooden frames, angled towards the
198 direction of the prevailing wind (Fig. S1d) with a maximum height of 1.8 m. The slight
199 reduction ($10 \pm 1\%$) in light transmission under the rain-out shelters (measured using a PAR
200 sensor; 400 to 750 nm; PP Systems International Inc., Amesbury, MA) was deemed unlikely
201 to impact significantly on grass growth, during the drought treatment. Rain was excluded
202 from the four drought and four flood+drought plots for 8 weeks during July to September
203 2016.

204 Recovery phase 2: All plots were monitored for two years following the end of the
205 drought period to determine the impact on ecosystem recovery in the absence of management
206 intervention. During this time, all plots were subjected to ambient conditions and livestock
207 and rabbits were excluded.

208

209 *2.3. Plot measurement frequency and ecosystem service assessment*

210 To determine the impact on ES, a range of above and below ground indicators were
211 identified which are directly linked to the provision of the six identified ecosystem services,
212 forage production, nutrient cycling, organic matter decomposition, climate regulation,
213 pollution regulation and biodiversity provision (Table 1; Adhikari and Hartemink, 2016;
214 Dominati et al., 2010; Rutgers et al., 2012; van Eekeren et al., 2010). Additionally, soil biota
215 play a key role regulating the ecosystem processes underpinning these ES (Creamer et al.,
216 2022). The importance of soil biodiversity in ecosystem multifunctionality is becoming
217 increasingly apparent (Creamer et al., 2022; Delgado-Baquerizo et al., 2020) along with the
218 role of plant-soil interactions (Forero et al.; Valencia et al., 2018). As such, the following
219 above and below ground biotic indicators supporting ecosystem processes were also
220 included: pasture community composition, microbial community characteristics (biomass and
221 PLFA), and earthworm abundance and biomass. Each plot was divided into dedicated areas
222 for monitoring (Fig. S1). During the flood, drought and recovery phases regular
223 measurements were taken from all plots (see Table S1).

224

225 *2.4. Above ground measurements*

226 Biomass was harvested from 40 × 40 cm quadrats within each plot on 5 occasions (1
227 month post flood, 1 week post drought, and at 6-, 12- and 24-months recovery). Plant material
228 was oven dried (80 °C, 24 h), weighed, ground (< 2 mm) and analysed for forage quality

229 including crude protein, metabolizable energy and digestibility (Sciante Analytical
230 Laboratories, York, UK). Annual biomass yield was calculated as tonnes per hectare for the
231 four treatments as the sum of the first 4 cuts for year 1 and as the 5th cut for year 2.

232 Plant surveys were undertaken over each entire plot by the same expert botanist in the
233 growing seasons throughout the duration of the experiment at the same time as the biomass
234 harvest. Presence and percentage cover (rounded to the nearest 1%) of all vascular plants and
235 bryophytes were recorded in each experimental plot. Cover of bare ground plus litter and
236 total bryophyte were also recorded and the species richness and Shannon-Wiener Diversity
237 were calculated (Magurran, 2013).

238 The percentage cover of each species was converted to a fraction. The proportion of
239 forage grasses which directly support livestock grazing and hence food production and
240 injurious weeds which detract from this provision (Maskell et al. 2020) were calculated along
241 with indicators of important species supporting biodiversity, namely; butterfly larval food
242 plants; supporting invertebrate populations (Lepidoptera) (Smart et al., 2000; Smart et al.,
243 2017); crop wild relatives as genetic insurance for food production (Jarvis et al., 2015) and
244 nectar-providing plants supporting pollinator biodiversity (Baude et al., 2016).

245

246 *2.5. Below ground measurements*

247 The plots were instrumented with SDI-12 soil moisture sensors (Acclima Inc.,
248 Meridian, ID) inserted horizontally at a depth of 5 cm, a static greenhouse gas chamber and
249 three MacroRhizon soil solution samplers (0.15 µm pore size; Rhizosphere Research
250 Products, Wageningen, The Netherlands) inserted at a 45° angle to a depth of 5 cm for
251 collection of soil pore water with two inserted within the main plot and one within the
252 greenhouse gas chamber. Soil sampling was undertaken immediately following the removal
253 of the flood or drought stress and at 6-, 12- and 24-month post drought recovery stages.

254 During each soil sampling event multiple soil cores ($n = 6$, $\phi = 1$ cm, depth = 0-10 cm) were
255 randomly taken from the indicated plot area (Fig. S1), homogenised and sieved (< 2 mm) for
256 analysis.

257 *Soil physical indicators:* An intact core (100 cm³, 0-10 cm depth) was taken from each
258 plot and weighed, dried (105 °C, 16 h), reweighed and dry bulk density and water filled pore
259 space determined.

260 *Soil (and floodwater) chemical indicators:* Soil moisture content was determined
261 gravimetrically (105 °C, 24 h). Soil pH and electrical conductivity (EC) were measured in 1:5
262 (w/v) soil to distilled water suspension using standard electrodes. Within 24 h of collection,
263 soils were extracted with 1:5 (w/v) soil-to-AcOH (0.5 M) and 1:5 (w/v) soil-to-K₂SO₄ (0.5
264 M). Macronutrient and trace element (P, K, Ca, Na, Al, Fe, Mg, Mn, Al) concentrations were
265 determined in the AcOH extracts via ICP analysis (Varian 720 ICP-OES). NO₃⁻ and NH₄⁺ in
266 the K₂SO₄ extracts were measured colorimetrically according to Miranda et al. (2001) and
267 Mulvaney (1996), respectively.

268 Total soil C and N were determined on oven-dried, ground soil using a TruSpec CN
269 Analyser (Leco Corp. St. Joseph, MI, USA). The tea-bag index (TBI) was used as a
270 standardised method to estimate both C decomposition (k) and C stabilisation (S) rate in soil
271 (Keuskamp et al., 2013). The difference between the two rates provides an indication of C
272 storage. Tea-bags were buried at 5 cm depth within each plot immediately prior to stress
273 initiation (Keuskamp et al., 2013) and recovered at the end of each stress period.

274 Soil redox measurements were taken on a weekly basis during the experiment with a
275 handheld SenTix[®] redox probe ($n = 3$ plot⁻¹; Wissenschaftlich-Technische Werkstätten
276 GmbH, Weilheim, Germany) inserted into the top 2 cm of soil. During the flood phase, soil
277 solution and the overlying floodwater was sampled weekly from the flood plots. Where
278 possible, the soil solution was sampled from the control treatment, however, there was only

279 sufficient soil moisture to extract a sample on limited sampling dates. These samples were
280 analysed for pH, EC, NH_4^+ and NO_3^- as described above and dissolved organic C (DOC)
281 using a Multi N/C 2100-S Analyser. Dissolved reactive P (DRP) and total dissolved P (TDP),
282 after acid persulfate digestion (Rowland and Haygarth, 1997) were determined using the
283 molybdate blue method of Watanabe and Olsen (1965). Dissolved organic P (DOP) was
284 calculated as the difference between TDP and DRP (AnalytikJena, Jena, Germany).

285 *Soil biological indicators:* Following collection, soil samples were immediately
286 frozen (-80 °C), freeze-dried and phospholipid fatty acid (PLFA) analysis undertaken
287 according to Bartelt-Ryser et al. (2005) to determine the soil microbial community profile.
288 Taxonomic groups were ascribed to individual PLFAs using the Sherlock[®] PLFA Method
289 and Tools Package (PLFAD1: Microbial ID Inc., Newark, DE) as outlined in Sánchez-
290 Rodríguez et al. (2019b).

291 Earthworm surveys were undertaken at the same timepoints as for soil sampling, by
292 excavating a soil pit (20 cm × 20 cm × 20 cm) and recovering live earthworms by hand.
293 Following soil excavation, 1 L of allyl isothiocyanate (1.3% v/v) in deionised water was
294 poured into the pit and left for 30 min to expel any deeper dwelling earthworms (Pelosi et al.,
295 2009). Collected earthworms were transported live to the laboratory in moist soil where
296 numbers of juvenile and mature earthworms were recorded and weighed.

297 *Greenhouse gas emissions:* A closed static chamber (40 cm × 40 cm × 25 cm) was
298 fitted to an aluminium frame installed in each plot at the beginning of the field experiment.
299 Immediately after closing each static chamber, a 20 ml gas sample was taken from the
300 headspace through a septum inserted in the lid using a needle attached to a 25 ml syringe at
301 time T0 and T60 mins. The gas samples were transferred to pre-evacuated glass vials (20 ml)
302 and analysed for CH_4 , CO_2 and N_2O using a Clarus 500 gas chromatograph equipped with a
303 HS-40 Turbomatrix autoanalyzer (PerkinElmer Inc., Waltham, MA).

304 Gas sampling was repeated 40 times between the start of the flood phase to day 235,
305 for each plot when no differences in gas flux had been observed between the treatment and
306 control plots for 90 days. The frequency of gas sampling was higher during the flood phase,
307 recovery phase 1 and the drought phase, and lower in recovery phase 2. At each gas sampling
308 time, soil temperature was recorded (0-2 cm depth) (T0 and T60 min) and the temperature
309 was used to correct the GHG fluxes. GHG fluxes were calculated using the difference in each
310 gas concentration between T0 and T60, the ratio between chamber volume and soil surface
311 area (Chadwick et al., 2014), and based on tests showing linearity (MacKenzie et al., 1998).
312 During the flood phase, an extension (25 cm height) was used to extend the static chamber in
313 the flood plots. Total cumulative fluxes were estimated using the trapezoidal rule (Rahman
314 and Forrester, 2021) for the different phases of the experiment; the flood phase (0 to 58 d),
315 the flood and recovery phase 1 (0 to 85 d), the drought phase (93 to 142 d), the drought and
316 recovery phase 2 (93 to 142 d), and the whole period of gas sampling (0 to 235 d). The total
317 GHG flux in CO₂ equivalents (kg C_{eq} ha⁻¹) was calculated by multiplying the total
318 cumulative fluxes of CH₄ by 28, CO₂ by 1, and N₂O by 265 (IPCC, 2013) and summing
319 them.

320

321 2.6. Data processing and statistical analyses

322 Soil moisture data from the Acclima moisture probes was converted to water filled pore
323 space (WFPS) as follows:

$$324 \quad WFPS (\%) = \frac{SWC}{\left(1 - \frac{BD}{PD}\right)} \times 100 \quad (\text{Eqn. 1})$$

325 where *SWC* = volumetric soil water content (vol. %), *BD* = soil bulk density (g cm⁻³) and *PD* =
326 particle density (2.65 g cm⁻³).

327 To examine the response of the above- and below-ground metrics to the treatments,
328 natural log response ratios (*lnRR*) were calculated as follows (Hedges et al., 1999):

329
$$\ln RR = \ln \left(\frac{X_{treatment}}{X_{control}} \right) \quad (\text{Eqn. 2})$$

330 where $X_{treatment}$ and $X_{control}$ are the mean values of each indicator under the treatment (flood,
331 drought and flood + drought) and the ambient control plots. In accordance with (Hedges et
332 al., 1999), we calculated the variance as follows:

333
$$\text{Variance} = (SD_{treatment}^2 / (nX_{treatment})) + (SD_{control}^2 / (nX_{control})) \quad (\text{Eqn. 3})$$

334 Subsequently the variance was converted to a 95 % confidence as shown below:

335
$$\text{Confidence interval} = 1.96 \times \sqrt{\text{Variance}} \quad (\text{Eqn. 4})$$

336 Finally, the response ratio and the corresponding confidence intervals were transformed to a
337 percentage change of the indicator from the control as follows:

338
$$(e^{\ln RR \text{ or Variance}} - 1) \times 100 \quad (\text{Eqn. 5})$$

339 The percent change and associated confidence interval was calculated for each above- and
340 below-ground indicator and the result plotted for three time periods: (1) immediately post
341 stress treatment; (2) after 1 year recovery; and (3) after 2 years recovery. If the 95 %
342 confidence interval did not overlap with zero, this indicated a significant change in the
343 indicator metric.

344 Changes in soil microbial communities were assessed by principal component analysis
345 (PCA) of PLFA taxonomic groups based on a data correlation matrix with principal
346 components (PCs). Analysis of variance (ANOVA) with Tukey's HSD post-hoc testing was
347 used to assess differences in the four treatments (for plant food production and biodiversity
348 provision plant indicators and greenhouse gas fluxes for each of the identified time periods).
349 The statistical cut-off for significance was $p < 0.05$. ANOVA and PCA were performed in the
350 statistical package SPSS software v22.0 (IBM Inc., Armonk, NY).

351

352 **3. Results**

353 *3.1. Assessment of treatment impacts*

354 Measurements of WFPS were used to validate the success of the treatments imposed
355 on the plots. At the initiation of flooding, WFPS increased from 34 % to a maximum of 85 %,
356 and remained at 79 ± 1 % for the duration of the flood period. The control plots fluctuated
357 between 20 and 45 % WFPS during the same period (Fig. S2). The drought treatments
358 resulted in a 71 % reduction in WFPS from 32 %, reaching 9 % by the end of the drought
359 period which was 63 % lower than the control. While the % reduction in WFPS over the
360 drought period was similar in the previously flooded plots, exhibiting a 70 % reduction from
361 54 % to 16 %, due to a higher initial value, this was only 47 % below the control. Both
362 flooding and drought exhibited a legacy effect from the stress, slowly converging toward the
363 control. The WFPS remained elevated in the flood plots and reduced in the drought plots,
364 compared to the control for the 6 months post drought when the sensors were removed. In
365 contrast, the flood+drought treatments showed a quicker recovery in WFPS following
366 drought treatment and WFPS was comparable to the control plots within 2 months (Fig. S2).

367 Flooding reduced the redox potential (Eh) from +297 mV (aerobic) to -92 mV
368 (anaerobic) after 50 d of flooding (Fig. S3), while in the control and drought plots Eh values
369 remained between +200 and +300 mV, throughout the trial. On removal of the flood water,
370 Eh rapidly increased to a maximum of 348 mV in the first 14 days and remained higher than
371 the control for 28 days before returning to the control range (+200 to +300 mV). While no
372 clear trends emerged during the drought phase all treatments plots showed a general trend of
373 higher Eh compared to the control during recovery 2 in the order of flood > flood+drought >
374 drought > control.

375 Neither flooding nor drought led to a change in soil temperature. In addition, no
376 unexpected periods of flooding or drought occurred during the experiment which might have
377 compromised the control treatments.

378

379 3.2. Above-ground indicator metrics

380 The response of each above-ground indicator immediately following the stress period
381 and following the three recovery periods are summarised in Figure 1. The annual above-
382 ground plant biomass for the flood plots was reduced by 30 % (CI: 24 %) compared to the
383 control, while the drought treatment did not significantly impact annual biomass production
384 (Fig. 1b). The reduction in annual biomass in flood+drought plots was greater than for flood-
385 only (54 % decrease, CI: 30 %). However, these changes in biomass were not carried over to
386 the following year (Fig. 1c). The change in pasture community which established post flood
387 (Table S2a) was accompanied by a change in the plant species supporting food provision and
388 biodiversity ES. The first month's pasture regrowth following flood stress showed a reduced
389 abundance of forage grasses, crop wild relatives (CWR) and plant species known to provide
390 food for butterfly lava (Fig. 2). There was also a slight increase in the number of tree and
391 shrub species seedlings (comprised of *Fraxinus excelsior*) and an increase in the number of
392 nectar-producing species. The reduction in forage grasses, CWR and butterfly lava food
393 remained for the first-year post flood, and the abundance of CWR and butterfly lava food
394 species remained reduced in the second year of recovery. Immediately following the drought,
395 the previously flooded plots continued to show a reduction in forage grasses, CWR and
396 butterfly lava food species. Additionally, these plots showed a small reduction in legumes.
397 The reduction in CWR and butterfly lava food species was still evident in the first year along
398 with an increase in the nectar species, but these changes were not present in the second year
399 of recovery. In contrast, post drought there was no change in these indicator species in
400 drought-only plots across the whole trial (Fig. 2, Table S2b).

401

402 3.3. Below-ground indicator metrics

403 The response of below-ground soil quality indicators to flooding and drought are
404 summarised in Figure 1. Overall, significant changes in a range of soil quality indicators were
405 apparent immediately post-stress removal, however, these changes progressively lessened
406 over the 2-year recovery period. In addition, the response of the combinatorial flood+drought
407 treatment was similar to the flood treatment.

408

409 3.3.1. Soil chemical indicators

410 Flooding reduced soil EC, soil available P (P_{acOH}), and C storage compared to the
411 control (Fig. 1a), and increased extractable NH_4^+ , Mn (Mn_{acOH}), and Fe (Fe_{acOH}). Drought
412 also reduced C storage and increased extractable NH_4^+ , NO_3^- , and Fe_{acOH} along with the soil
413 EC. The flood+drought reduced available P (P_{acOH}) and increased soil EC, Mn_{acOH} , Fe_{acOH} ,
414 NH_4^+ and NO_3^- .

415 During the first year of recovery P_{acOH} was reduced in all stress combinations (Fig.
416 1b) but was increased compared to the control in the second-year post stress in the flood and
417 flood+drought plots (Fig. 1c). Extractable Fe_{acOH} and Mn_{acOH} concentrations were increased
418 in the flood plots and Mn_{acOH} increased in the flood+drought plots during the first year. Two
419 years post stress Fe_{acOH} concentrations were reduced in both treatments. Flooding increased
420 soil NH_4^+ , and drought increased both soil NH_4^+ and NO_3^- but this did not persist through the
421 recovery phase. However, at the two-year recovery point NO_3^- concentrations were increased
422 in the flood plots.

423 Due to low moisture conditions in some treatments, soil solution samples were only
424 consistently collectable from the flood plots, and only during the flood phase and recovery
425 phase 1. There was a general trend of increasing DOC and decreasing DRP and NO_3^- as the
426 flood period progressed (Fig. 3). On removal of the flood water there was a spike in DOC
427 concentrations accompanied by a spike in DOP followed by a sharp decline. Additionally,

428 there was a large spike in the NH_4^+ concentrations 13 d following removal of the flood water
429 to a maximum of 131 mg N L^{-1} which quickly declined to background levels (0.04 mg N L^{-1})
430 by day 21. A sharp spike in NO_3^- concentrations to a maximum concentration of 34.9 mg N
431 L^{-1} was observed 7 days after flood removal followed by a more gradual decline to
432 background levels (0.83 mg N L^{-1}) by day 27.

433

434 3.3.2. *Soil biological indicators*

435 Distinct changes in microbial community structure were observed with respect to both
436 sampling time and treatment. Seasonality played a key role in controlling the distribution of
437 PLFA markers with marked separation between the stress and recovery periods and a clear
438 shift towards more Gram-positive dominated communities as time progressed. However,
439 treatment effects on community structure were also observed (Fig. 4). Flood stress increased
440 the dominance of Gram-negative bacteria while drought led to a shift towards actinomycete
441 markers. The flood and flood+drought plots also had a reduction in the fungi-to-bacteria ratio
442 and amount of arbuscular mycorrhizal fungi (AMF) in the first- and second-year post stress
443 (Fig 1.b.c). The PLFA markers within the flood plots showed a larger separation from the
444 control treatments compared to the drought and flood+drought plots after the first and second
445 year of recovery. After 1-year of recovery there were three distinct groups of PLFA
446 distributions, (1) control, (2) drought and flood+drought, and (3) flood. At the 2-year
447 recovery point the control group remained separated from the three stress treatments (Fig. 4).
448 The size of the microbial biomass was also slightly increased in the flood and drought plots
449 during the first year and in the flood-only and flood+drought plots in the second year relative
450 to the control. Earthworm abundance increased following removal of flooding, however, it
451 decreased following the removal of drought (Fig. 1).

452

453 3.3.3. Greenhouse gas emissions

454 During the flood phase, CO₂ emissions were higher in the control plots in comparison
455 with the flood plots (up to 300 mg C m⁻² h⁻¹ vs. <1 mg C m⁻² h⁻¹, respectively; Fig. 5b).
456 During the drought CO₂ emissions were similar for the control and flood plots and reduced in
457 the drought and flood+drought plots (between 80–320 mg C m⁻² h⁻¹ vs. 30–190 mg C m⁻² h⁻¹,
458 respectively; Fig. 5b). During recovery phase 1 and 2, the CO₂ emissions were similar
459 between all treatments. Overall, flood or drought caused a large decrease in cumulative CO₂
460 emissions while their combination (flood+drought) resulted in the lowest cumulative
461 emissions (Table 2).

462 No N₂O emissions were observed during the flood phase for the control or flood plots.
463 During the drought phase, application of N fertiliser resulted in a peak in N₂O emissions in
464 the control and flood plots (up to 0.43 and 0.82 mg N m⁻² h⁻¹, respectively) while no
465 emissions were observed in the drought or flood+drought plots (<0.07 mg N m⁻² h⁻¹). The
466 spring flood caused N₂O peaks of 0.2–1.2 mg N m⁻² h⁻¹ (60–85 d after flooding) that were
467 more frequent and higher than the fluxes observed following N fertilization in the control
468 (Fig. 5c). No N₂O emissions were observed in any of the plots in the second recovery phase
469 (Fig. 5c). The highest total cumulative N₂O emissions were calculated for the flood+drought
470 plots (4.16 kg N ha⁻¹), followed by flood plots (3.35 kg N ha⁻¹), control plots (1.01 kg N
471 ha⁻¹) and drought plots (0.41 kg N ha⁻¹); the differences were significant between the
472 flood+drought and the drought plots (Table 2). These emissions occurred mainly during
473 recovery phase 1 (Table S3). The expected increase in N₂O emissions due to N fertilisation
474 (during the drought phase) was influenced by the treatment where total cumulative N₂O
475 emissions followed the order flood ≥ control = flood+drought ≥ drought (Table S3).

476 CH₄ emissions remained close to zero except immediately after flood water removal
477 (58 d after flooding) when they peaked (1–2 mg C m⁻² h⁻¹) in the flooded plots (Fig. 5a).

478 Mean cumulative CH₄ fluxes were small and not significant between treatments over the 2-
479 year experiment (Table 2). Overall, the control and drought treatments acted as a sink for
480 CH₄ (-0.38 and -0.20 kg C ha⁻¹, respectively) while the flood and flood+drought were a
481 minor source of CH₄ (0.98 and 1.49 kg C ha⁻¹, respectively). Table S3 shows that (i) the
482 control plots were a sink during most of the experiment (except in the second recovery
483 phase), (ii) flood induced CH₄ emissions (mainly immediately following flood water
484 removal), and (iii) drought reduced CH₄ emissions (only during the drought phase).

485 When considering the combined emissions of GHGs in terms of total CO₂-
486 equivalents, significant differences were apparent and followed the series control ≥ flood ≥
487 drought = flood+drought (Table 2). If direct CO₂ emissions are excluded, the pattern was
488 similar to total cumulative N₂O emissions (i.e. flood+drought ≥ flood ≥ control ≥ drought)
489 over the 2-year period.

490

491 **4. Discussion**

492 This study utilised a suite of above and below ground properties of an improved sheep
493 grazed pasture to determine the impact of both single extreme events characterised by
494 contrasting rainfall patters (flood or drought), and the combined impact of a swift transition
495 from flood to drought conditions on ecosystem functioning. Our results clearly show three
496 key findings: (1) this grassland system was more resistant and resilient to drought conditions
497 than flooding, (2) the combination of flood and drought did not magnify the impact of the
498 single stressors with some indication of faster recovery times and (3) persistent changes in
499 plant and microbial communities post flood likely have implications for future ecosystem
500 service provision.

501

502 *4.1. Impact of contrasting single extreme events (i.e. flood vs. drought).*

503 Both flooding and drought were found to have a large impact on the above and below
504 ground indicators of ecosystem function in improved grasslands supporting out hypothesis
505 that extreme weather events will negatively impact ecosystem processes. However, flooding
506 was more detrimental with 17 indicators being different to the control immediately post stress
507 compared to 8 indicators for drought (Fig. 1). Flood plots also showed a difference in 7
508 indicators two years after the stress compared to just 1 indicator two years post drought.
509 Together these results show lower resistance and resilience to flooding than drought.
510 Further evidence of this was found in the shifts in microbial community, as indicated by
511 PLFA markers. Larger shifts occurred for the flood plots and remained more distinct for
512 longer compared to drought. This prolonged shift in microbial communities agrees with
513 previous mesocosm studies, where a shift in PLFA markers occurred following extreme
514 flooding (80 days) and persisted for the length of the monitored recovery period of 60 days
515 (Sánchez-Rodríguez et al., 2018). In contrast, microbial communities in grasslands have been
516 shown to be largely resistant to drought (Birkhofer et al., 2021; Blankinship et al., 2011) and
517 unlike for flooding there was no change in microbial biomass during the drought period
518 Drought did induce a larger shift towards actinomycete taxa which have been shown to be
519 more resistant to dry and wetting cycles (Acosta-Martínez et al., 2014; Ochoa-Hueso et al.,
520 2020; Ochoa-Hueso et al., 2018). However, seasonality had a large control on PLFA marker
521 distribution likely due to changing environmental conditions. Interestingly, the shift towards
522 actinomycetes compared to the control also occurred in the flood plots post drought, despite
523 no drought treatment being imposed. This indicates that immediately following
524 environmental stress, microbial communities are more responsive to changes in
525 environmental conditions compared to unstressed communities demonstrating lower
526 resilience.

527 Environmental stressors can have detrimental impacts on earthworm communities
528 (Singh et al., 2020). Earthworms are considered keystone species within soil environments
529 due to their influence on soil chemical, physical and biological processes (Blouin et al., 2013;
530 Jones et al., 1994). Generally, prolonged flooding has been shown to reduce earthworm
531 abundance across different ecosystems (Singh et al., 2020). However, our study showed a 76
532 % increase in earthworm numbers (predominantly juveniles) immediately following removal
533 of the overlying floodwater. Visual observation suggested this was the result of recent
534 hatching, with some earthworms still attached to cocoons. This finding has only been
535 reported for one study in annually flooded Alder forests (Emets, 2018). We hypothesise that a
536 sudden change in osmotic potential ruptured the cocoons outer shell leading to premature
537 emergence. In contrast drought did not impact on earthworm biomass or abundance, likely
538 due to their ability to enter estivation in response to low moisture conditions, indication of
539 which was observed in the post drought sampling with many individuals exhibiting
540 characteristic knot configurations (McDaniel et al., 2013). The trend post flood did not persist
541 into the recovery phase once cocoons within the control soils hatched, and under the single
542 flood stress condition this short-lived increase in earthworm numbers is unlikely to influence
543 overall ecosystem function.

544 Above and below-ground biodiversity has been shown to be positively related to
545 ecosystem multifunctionality (Delgado-Baquerizo et al., 2020) and even small changes in
546 microbial structure (such as community composition shifts) can have a large impact on
547 function due to altered metabolic capability of the ecosystem (Crowther et al., 2019). The
548 pronounced and persistent change microbial communities induced by flooding were also
549 accompanied by increased plant diversity and richness which will influence soil function with
550 knock-on effects on the provision of ecosystem services while the impact of drought appears
551 to be minimal and immediate with limited longer-term impacts. Where significant indicator

552 responses to treatments were found, their impact on the 6 highlighted ES will be discussed in
553 section 4.3.

554

555 *4.2. Impact of consecutive contrasting extreme event types (i.e. flood vs. drought).*

556 The combination of flood and drought stress did not increase the number of indicators
557 affected compared to flood alone either immediately post stress (15 for flood+drought cf. 17
558 for flood) or at the 2-year recovery (5 flood+drought cf. 6 for flood). In fact, there was some
559 indication within the microbial community of faster recovery in community composition
560 evidenced by reduced separation of PLFA markers from the control plots 1 year post stress
561 compared to flood alone. We suggest that the more rapid reduction in soil moisture post flood
562 in the flood+drought plots may have increased the resilience of microbial population in line
563 with the first of the two alternative hypotheses that subsequent events of contrasting typology
564 will promote ecosystem recovery (hypothesis A).

565 However, while hypothesis A may fit the microbial community response the impact
566 on earthworm populations showed contrasting results. The large increase in juvenile
567 earthworm numbers post flood was followed by a dramatic decrease of 85 % in abundance
568 compared to the control post drought despite no significant change in the drought plots. We
569 suggest that the premature emergence of juveniles during flooding may have implications for
570 the resilience of the community if a second extreme event occurs prior to this community
571 reaching maturity, in line with the second alternative hypothesis B. Larger soil invertebrates
572 including earthworms are especially important for maintaining optimal ecosystem functions
573 (Delgado-Baquerizo et al., 2020) and loss of earthworm populations could have large-scale
574 consequences for ecosystem service provision.

575

576 *4.3 Implications for ecosystem service provision*

577 Flood, drought and the combined flood+drought had an impact on both the indicators directly
578 related to the 6 specific ecosystem services considered.

579 Forage production: The reduction in annual biomass and the slow recovery of forage grass
580 species following flooding will likely decrease production at least during the event year.
581 Additionally, flooding resulted in large periods of increased bare ground while the plant
582 community re-established itself. The extremes of wet and dry conditions will limit the pool of
583 successful colonists. A likely scenario in the longer term is larger areas of bare ground with
584 varying cover of transient and persistent weed species. Recolonisation of bare ground with
585 invasive weed species is a concern to farmers (ADAS, 2014). However, in our study neither
586 flooding nor drought or the combination led to a change in the % cover of injurious weeds
587 compared to ambient conditions.

588 These results indicate that flooding will lead to a reduction in forage provision and
589 subsequent production metrics at least during the stress year while drought may have limited
590 impact in this grassland. Farmers may need accept reductions in profit margins to bring in
591 additional forage to meet feed demands, increasing production costs or reduce stocking rates.

592 Nutrient cycling: The impact of extreme weather events on soil fertility was evidenced
593 through immediate and prolonged changes in extractable soil nutrients. While soil P
594 concentrations were largely unaffected by drought stress, flooding altered soil P in different
595 directions at the various timescales. The reduction in plant-available P post flood is consistent
596 with the reductive dissolution of P bound to iron or manganese oxides within the clay fraction
597 of the soil under the observed negative redox potentials (Sánchez-Rodríguez et al., 2019a;
598 Sánchez-Rodríguez et al., 2019b). The P released from the clay surfaces is prone to leaching
599 while the iron and manganese oxides are quickly precipitated within aerobic micro-sites
600 leading to the accumulation of Fe_{acOH} and Mn_{acOH} . A general downward trend in soil solution
601 DRP as the flood period progressed supports this theory. This reduction in P_{acOH} persisted for

602 the first year of recovery potentially impacting sward recovery. Hence, farmers should be
603 encouraged to test for soil P status following prolonged flooding, as this may have changed.

604 N dynamics in soils are tightly coupled to soil moisture (Bowles et al., 2018).
605 Consequently, in contrast to P, soil mineral N concentrations were influenced by both
606 drought and flood stress. Suppression of plant growth combined with low soil moisture and
607 limited leaching potential during drought led to the accumulation of NH_4^+ and NO_3^- in the
608 soil. While the accumulation of NH_4^+ during flood stress was accompanied by no change in
609 NO_3^- concentration despite reduced plant N requirements. We ascribe the greater
610 accumulation of NH_4^+ during the flood phase compared to the drought to a combination of
611 (1) suppressed microbial nitrification at low Eh potentials during flooding, and (2) increased
612 mineralisation of organic matter introduced as dead plant material. Waterlogged conditions
613 increase the leaching potential of NO_3^- , and the rate of denitrification (Rohe et al., 2021)
614 explaining the lack of soil NO_3^- response. The accumulation of mineral N immediately post
615 drought or flood stress provides essential nutrients for plant growth following removal of the
616 stress and is essential to the resilience of pasture systems (Oram et al., 2020). This likely
617 contributed to the lack of annual yield loss seen post drought but was not sufficient to
618 mitigate the within-event year impact of flooding on annual biomass.

619 Drought did not influence soil fertility metrics in the longer term. However, differing
620 P_{acOH} and NO_3^- responses post flood and during the recovery periods suggest a longer-term
621 impact on N and P dynamics in the flood plots. During recovery phase 2, no fertiliser was
622 applied, and livestock were excluded, therefore the increase in P_{acOH} and NO_3^- concentrations
623 in the flood plots suggests increased biogeochemical cycling of nutrients to replace those lost
624 from the system. This recovery of soil fertility in the flooded plots was further evidenced by a
625 recovery of annual biomass to that of the control plots in year 2.

626 The exact mechanisms for these patterns of soil fertility are unclear but may be due to
627 a change in soil-plant-microbe interactions. In the absence of fertiliser inputs, soil fertility is
628 largely driven by microbial processing of organic matter (Risch et al., 2019). Manipulation of
629 rainfall has been shown to alter soil microbial community structure (Ochoa-Hueso et al.,
630 2020) and microbial activity via increased extracellular enzyme expression (Ochoa-Hueso et
631 al., 2018). The higher microbial biomass in flooded soils 1 and 2 years post stress can be
632 attributed to increased inputs of dead organic matter and an associated resource pulse post
633 flooding (Wright et al., 2015). Furthermore, plant-soil feedback mechanisms are key drivers
634 of ecosystem processes and plant associated microbiomes are profoundly influenced by
635 abiotic factors and changing environmental conditions (Pugnaire et al., 2019; Saijo and Loo,
636 2020). The shift in plant and microbial community composition observed during pasture
637 regrowth may have influenced the soil chemistry and the root associated microbiome, as
638 evidenced by the shift in PLFA markers, including markers for AMF, with implications for
639 nutrient acquisition strategies (Lozano et al., 2021). Further work, on the detailed changes in
640 microbial communities, especially gene expression of N and P cycling and investigation of
641 the changes in extracellular enzyme expression may further elucidate the mechanisms
642 involved.

643 Organic matter decomposition: Results from the tea-bag index found reduction in the C
644 stabilisation rate during both the flood and drought phase, and a reduction in C
645 decomposition rate during the flood with impacts on below ground C storage.

646 Anoxic conditions are generally considered to decrease C decomposition and favour
647 C accumulation (Greenwood, 1961; Reichstein et al., 2013). However, while the
648 decomposition rate decreased during flooding, C storage also decreased during both the flood
649 and the drought phases. The reduction in C storage was driven by a reduction in C
650 stabilisation rates, likely due to reduced rhizodeposition. Drought has been shown to alter

651 allocation of C within plants (Bahn et al., 2013) and affect exudate quantity and quality
652 (Williams and de Vries, 2020) but the impact of flooding on these dynamics remains a
653 significant gap in the literature. Additionally, iron redox cycles during flooding can release
654 mineral protected C with leaching of DOC via the soil solution. Measurable changes in soil C
655 stores take time to develop. However, the persistent shift in microbial composition post flood,
656 could impact on C turnover especially if more resource-conservative stress tolerant traits are
657 favoured over enzyme expression (Wang and Allison, 2021).

658 Climate regulation: Soil plays an important role in regulating the climate and can act both as
659 a sink for C and a source of GHGs (Lal et al., 2021). Flood and drought decreased CO₂
660 emissions and cumulative CO₂ fluxes during the stress period due to a reduction in soil
661 microbial and root respiration with the largest effect on total flux seen for the combined
662 events. However, the lower cumulative CO₂ emissions during drought in the previously
663 flooded plots indicates that the negative impact of the combination of stresses on soil was not
664 additive.

665 Flooding has a significant impact on the production of CH₄ due to the generation of
666 negative redox potentials (Zhang and Furman, 2021). Suppressed gas exchange under
667 waterlogged conditions delayed the pulse of CH₄ release until the removal of the flood water.
668 In the flooded soils, the drought period produced a switch from a CH₄ source to a CH₄ sink.
669 The sustained release of CH₄ from the flooded soils is likely due to the lag in WFPS
670 recovery, but changes in the microbial community structure may also play a role.

671 Nitrous oxide makes an important contribution to GHG emissions from grazed
672 pasture systems with soil moisture, N fertiliser application (Cardenas et al., 2019) and excreta
673 deposition (Chadwick et al., 2018) being the main drivers of N dynamics. Flood stress had
674 the largest impact on N₂O emissions with a large pulse emitted once the floodwater was
675 removed and a second pulse associated with fertiliser application. During the flood and the

676 first days of recovery, NH_4^+ is the predominant inorganic N form within the soil mineral and
677 solution phases as anaerobic conditions suppress nitrification. Subsequent mineralization of
678 dead vegetation, mesofauna and soil microorganisms resulted in a pulse of C and N release
679 facilitated by favourable summer temperatures (Kirwan and Blum, 2011; Sánchez-Rodríguez
680 et al., 2019b). This resulted in a peak in NH_4^+ during soil recovery. Concurrent nitrification
681 and denitrification occurring at microsites within the soil led to the observed N_2O emissions
682 in the soil recovery phase after flooding (Miniotti et al., 2016; Zhu et al., 2013) and
683 contributed to the observed decrease in NH_4^+ and NO_3^- , in the soil solution. Additionally,
684 NH_3 volatilization (when the soil was flooded and saturated) would have reduced the NH_4^+
685 concentration in the soil solution (Verhoeven et al., 2018). During longer-term recovery,
686 increased N demand due to a resumption of the growth of plants (seed bank) and soil
687 microorganisms under non-anoxic conditions led to a gradual decline in NO_3^- in the soil
688 solution, in line with N_2O peaks. However, this reduction was slower than for NH_4^+ , which
689 indicates that the gradual recovery of plants and microorganisms which was insufficient to
690 fully utilise the available N resource. Current understanding of N dynamics in agricultural
691 systems identifies N fertilisation as the main driver of spikes in N_2O emissions. However, the
692 N_2O pulse following floodwater recession was greater than that produced by N fertilisation.
693 Consequently, flooding caused the highest GWP (excluding CO_2) in line with Hou et al.
694 (2000), especially in combination with drought. These data can help to fill knowledge gaps
695 related to GHG emissions from agricultural soils (Bianchi et al., 2021). Furthermore,
696 alteration of the soil microbial communities is likely to play a large role through the alteration
697 of N dynamics and the long-term effect of flood induced shifts in microbial structure on GHG
698 emissions requires further investigation.

699 Pollution regulation: The risk of erosion and runoff generation immediately post flood is
700 increased due to the severe reduction in plant cover and the high WFPS, reducing infiltration

701 capacity of the soil, and remains a threat for the first year of recovery along with an increased
702 risk of nutrient and sediment transport. Conversely, 8 weeks of drought did not increase the
703 proportion of bare ground present retaining soil protection to future heavy rain events,
704 minimising the erosion risk and impact on water quality.

705 Flood events have been shown to export significant amounts of P to receiving
706 waterbodies (Ockenden et al., 2017). Traditionally, P leaching has not been considered to be
707 a major transport pathway in clay-rich soils. However, evidence of P leaching to groundwater
708 in certain soil, hydrologic and management conditions is growing (McDowell et al., 2019;
709 Stuart and Lapworth, 2016) and in flooded soils this may be enhanced by reductive
710 dissolution cycles of Fe/Mn oxides. The general downward trend in DRP over the flood
711 period indicates loss of P to leaching while the spikes in concentration are likely due to
712 localised release of P from iron oxides in response to anoxic conditions developing. Further
713 evidence of loss of P from the soil is seen through the reduction in P_{acOH} post flood. Unlike P,
714 NO_3^- is prone to leaching and the general downward trend in soil solution NO_3^- , coupled with
715 low N_2O emissions during the flood phase suggest possible loss to groundwater. During this
716 period, the concentrations detected were low ($< 0.25 \text{ mg } NO_3^- \text{ L}^{-1}$) and unlikely to pose a
717 significant water quality risk. However, immediately post flood there is an increased risk of N
718 and P leaching due resumption of microbial activity leading to mineralisation of dead plant
719 material and nitrification processes converting accumulated NH_4^+ to NO_3^- , evidenced by a
720 large, and environmentally significant, spike in DOC, DOP and NO_3^- concentrations within
721 the soil solution followed by a sharp decline.

722 Biodiversity provision: Grasslands provide important floral resources for pollinating species.
723 However, the low biodiversity within many agricultural pastures supporting livestock
724 production is contributing to global declines in species (Sánchez-Bayo and Wyckhuys, 2019).
725 While drought had no impact on plant community composition, the increase in diversity and

726 richness recolonising the flood plots post stress may promote greater insect populations
727 including pollinating species evidenced by a persistent increase in nectar-containing species
728 in the flood plus drought treatment. However, in contrast grassland crop wild relatives are
729 largely composed of flooding-intolerant grass species (Jarvis et al., 2015) and flooding
730 reduced their abundance for 1 year post-flood. Crop wild relatives contribute towards an
731 ecosystem's genetic resource provision (Jarvis et al., 2015) and often exhibit useful traits
732 related to stress tolerance (Castañeda-Álvarez et al., 2016). However, if flooding becomes
733 more widespread, reduction of this valuable resource could have negative impacts for
734 biodiversity provision and future agricultural plant breeding programs aimed at producing
735 plant varieties more resilient to climate change.

736

737 **5. Conclusions**

738 Extreme weather events are increasing in frequency and magnitude. Here we show that
739 flooding and drought have a significant impact on a range of above and below ground
740 indicators of ecosystem function and that this pasture system is more resistant and resilient to
741 drought than to flood. Contrary to our prediction, the compound flood+drought treatment did
742 not further exacerbate the flood impact. In fact, there was some indication of more resilience
743 in soil and microbial parameters suggesting contrasting event types may promote recovery
744 through a more rapid return to ambient soil moisture conditions. The immediate impact of
745 flooding on all ecosystem services was negative, especially for production and climate and
746 water regulation. Flooding stress caused pronounced and persistent shifts in soil microbial
747 and plant communities with large implications for nutrient cycling and long-term ES
748 provision. However, determination of the mechanisms behind these changes and the impact
749 of these shifts along with the complex plant-soil-microbe interactions, on whole ecosystem
750 function remains a key research priority. The observed increase in soil nutrient concentrations

751 immediately post flood and drought stress has the potential to provide a resource pulse for
752 pasture growth and recolonization. Soil testing is essential when determining fertiliser rates to
753 minimise nutrient loss and minimise impacts on water quality and GHG emissions.
754 Furthermore, the large pulse of N₂O emissions is striking and with the projected increase in
755 the frequency, severity and spatial extent of flood events further research is required to factor
756 flood-induced N₂O pulses into GHG emission projections.

757

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766

767 **Author contribution**

768 Davey L. Jones, David R. Chadwick, Dylan Gwynn-Jones, Rosalind J. Dodd and Paul. W.
769 Hill conceived the study and set up the experiment. Rosalind J. Dodd, Antonio R. Sánchez-
770 Rodríguez and Simon M. Smart undertook the main fieldwork, laboratory studies and data
771 analysis. Rosalind J. Dodd wrote the first draft of the manuscript and all authors contributed
772 substantially to revisions and have given final approval of the submitted manuscript.

773

774 **Conflict of interest**

775 The authors declare no competing interests.

776

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1121 **Figure legends**

1122 **Fig. 1.** Response of above- and below-ground indicators (a) immediately post flood or
1123 drought stress, (b) 1 year and (c) 2 years following the stress event. Effect sizes represent
1124 variation from the control calculated from the LnR/R response ratio as outlined in the
1125 methods. Error bars represent the 95 % confidence interval. Open symbols represent
1126 significant effects (95 % CI not overlapping Zero). The number of indicators significantly
1127 different to the control at the three time-points for the three treatments are shown in (d). In all
1128 panels AMF denotes arbuscular mycorrhizal fungi.

1129

1130 **Fig. 2.** Ecosystem service indicator scores provided by the pasture community (a) at the 1-
1131 month pasture regrowth stage post flood, (b) immediately post drought, (c) 6-months post
1132 drought, (d) 1-year post drought and (e) 2-years post drought. The letters denote significant
1133 differences between treatments at the $p < 0.05$ as determined by the Tukey HSD test of
1134 multiple comparisons and *n.s.* denotes no significant difference.

1135

1136 **Fig. 3.** Mean temporal dynamics of soil solution (a) dissolved organic carbon (DOC), (b)
1137 dissolved reactive phosphorus (DRP), (c) dissolved organic phosphorus (DOP), (d)
1138 ammonium (NH_4^+) and (e) nitrate (NO_3^-) within the flood plots during the flood phase and
1139 recovery phase 1. Error bars show the standard error of the mean.

1140

1141 **Fig. 4.** Principle component analysis for PLFA (a) in the four treatments across the five
1142 sampling dates, immediately post flood, immediately post drought and 6, 12 and 24 months
1143 post drought and (b) vectors indicate the direction of shift towards specific taxonomic groups
1144 85 % of the variance in PLFA markers was explained by PC1 (x-axis) and a further 10 % was

1145 explained by PC2 (y-axis). Ellipses group the responses to the specific time period and errors
1146 bars show the standard error of the mean for the four treatments.

1147

1148 **Fig. 5.** Greenhouse gas emissions for each treatment showing hourly CH₄ (a-d), CO₂ (d-h)
1149 and N₂O (i-l) fluxes (mean ± standard error, *n* = 4) in each period of the experiment (flood
1150 phase, recovery phase 1, drought phase and recovery phase 2) for the different treatments
1151 assessed in this study (control, flood, drought and flood + drought). The arrow and F indicate
1152 a fertilization event.

1153

1154

1155 Table 1. Indicator metrics associated with the 6 selected provisioning, supporting and regulating
 1156 services provided by agricultural grasslands.

| Environment Service category | Service | Indicators | Measurement |
|------------------------------|------------------------------|--|---|
| Provisioning | Food provisioning | Biomass, forage grass cover, weed species cover | Pasture cut, plant surveys |
| | Biodiversity provision | Plant species richness + diversity, nectar producing species, butterfly larvae food species, crop wild relatives | Plant surveys |
| Regulating | Climate regulation | CO ₂ , CH ₄ , N ₂ O emissions | Static chamber greenhouse gas emission measurements |
| | Pollution regulation | % ground cover, soil solution chemistry (DOC, NO ₃ ⁻ , NH ₄ ⁺ , DRP, DOP) ¹ | Plant surveys, rhizon sampling of soil solution (during flood phase only) |
| Supporting | Nutrient cycling | Soil extractable NO ₃ ⁻ , NH ₄ ⁺ , P | Soil samples |
| | Organic matter decomposition | C stabilisation rate (<i>S</i>), C decomposition rate (<i>k</i>) | Tea-bag index |

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1158 ¹ DRP = dissolved reactive phosphorus, DOC = dissolved organic phosphorus, DOC = dissolved organic
 1159 carbon.

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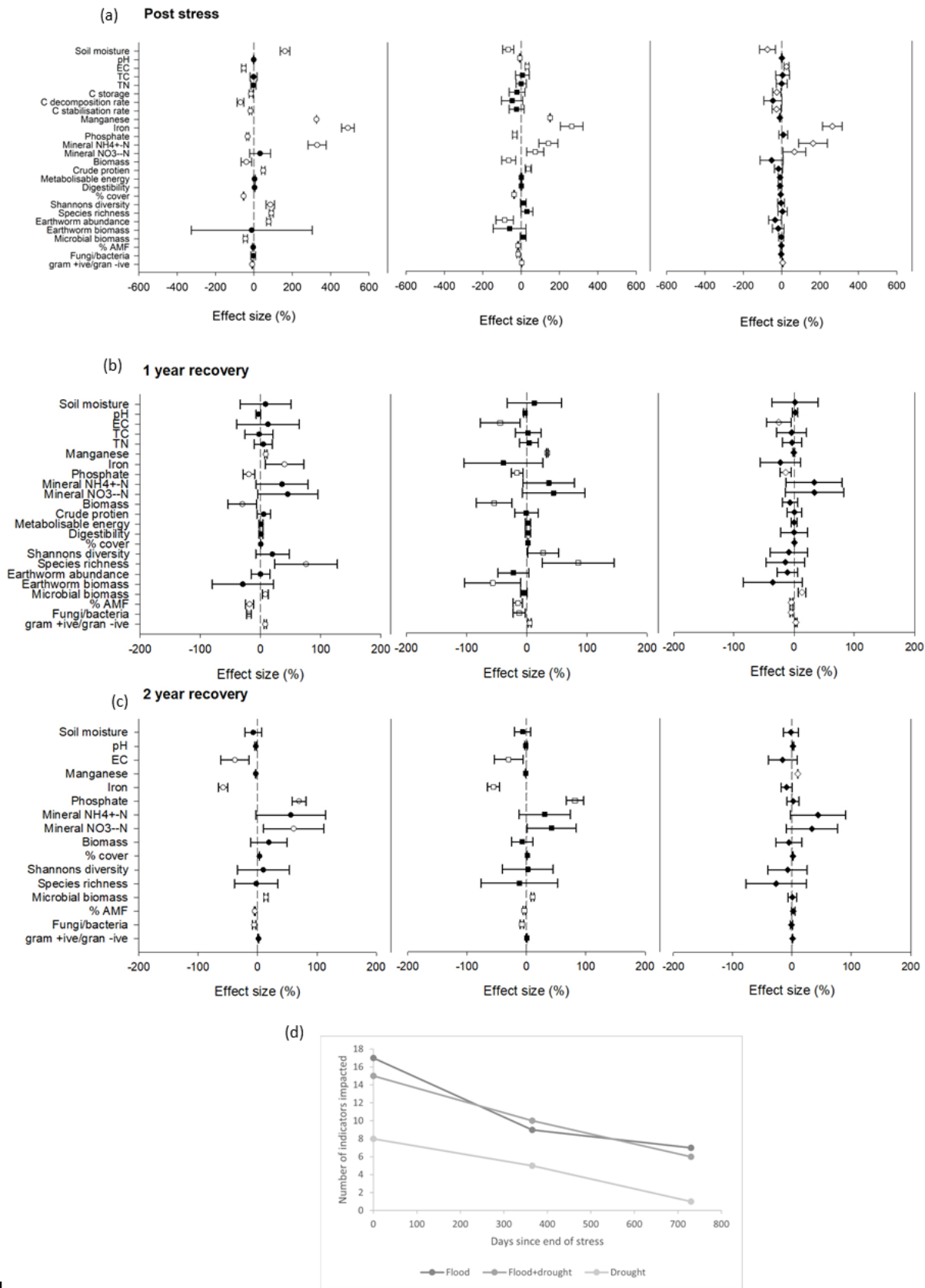
1161 Table 2. Mean cumulative GHG fluxes and total CO₂ equivalents with and without CO₂ emissions
 1162 along with the standard error. The letters denote significant differences between treatments
 1163 according to the Tukey's HSD test at the *p* < 0.05 level of significance.

| Treatment | CH ₄ (kg C ha ⁻¹) | CO ₂ (kg C ha ⁻¹) | N ₂ O (kg N ha ⁻¹) | GWP (exc. CO ₂) (kg CO ₂ eq ha ⁻¹) | GWP (inc. CO ₂) (kg CO ₂ eq ha ⁻¹) |
|--------------------|---|---|--|--|--|
| Control | -0.38 ± 0.18 a | 8111 ± 192 a | 1.01 ± 0.14 ab | 273 ± 40 ab | 7915 ± 274 a |
| Flood | 0.98 ± 0.72 a | 4651 ± 1213 ab | 3.35 ± 1.03 ab | 874 ± 264 ab | 5238 ± 1347 a |
| Drought | -0.20 ± 0.03 a | 4726 ± 1059 ab | 0.41 ± 0.06 b | 108 ± 15 b | 4602 ± 1062 a |
| Flood & drought | 1.49 ± 1.16 a | 2942 ± 617 b | 4.16 ± 1.43 a | 1060 ± 343 a | 3813 ± 925 a |
| <i>p</i> value | 0.215 | 0.009 | 0.031 | 0.025 | 0.059 |

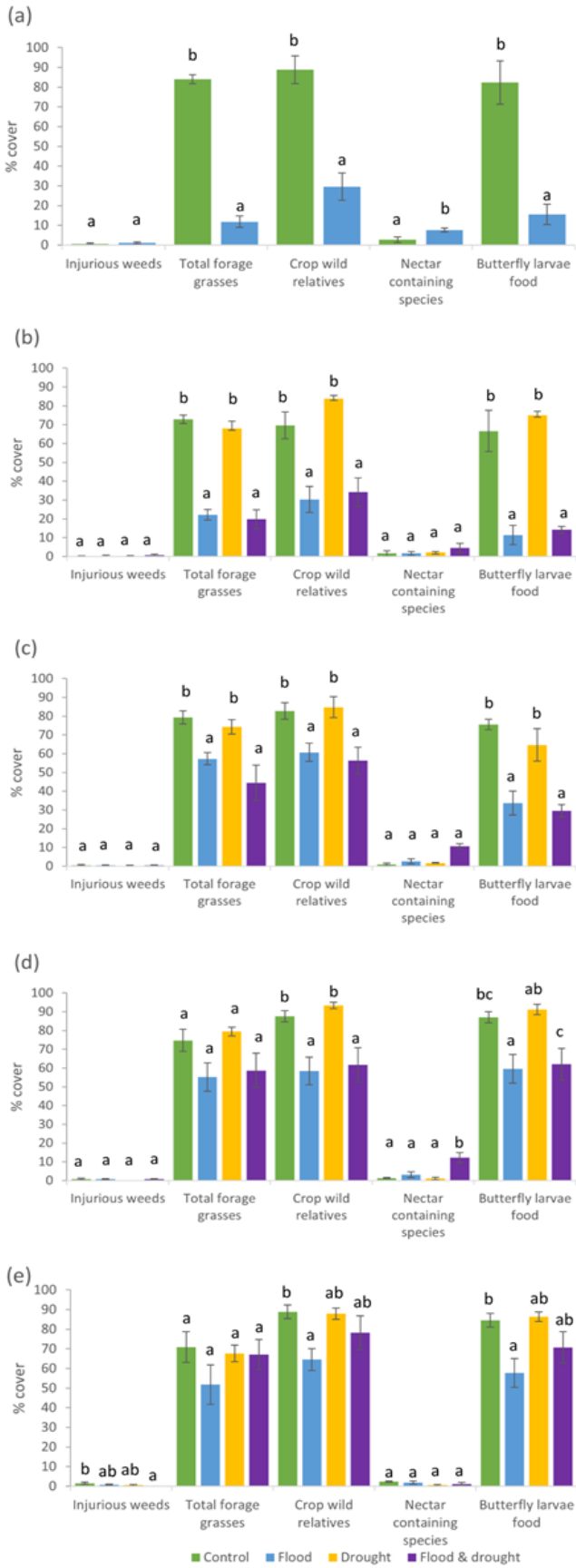
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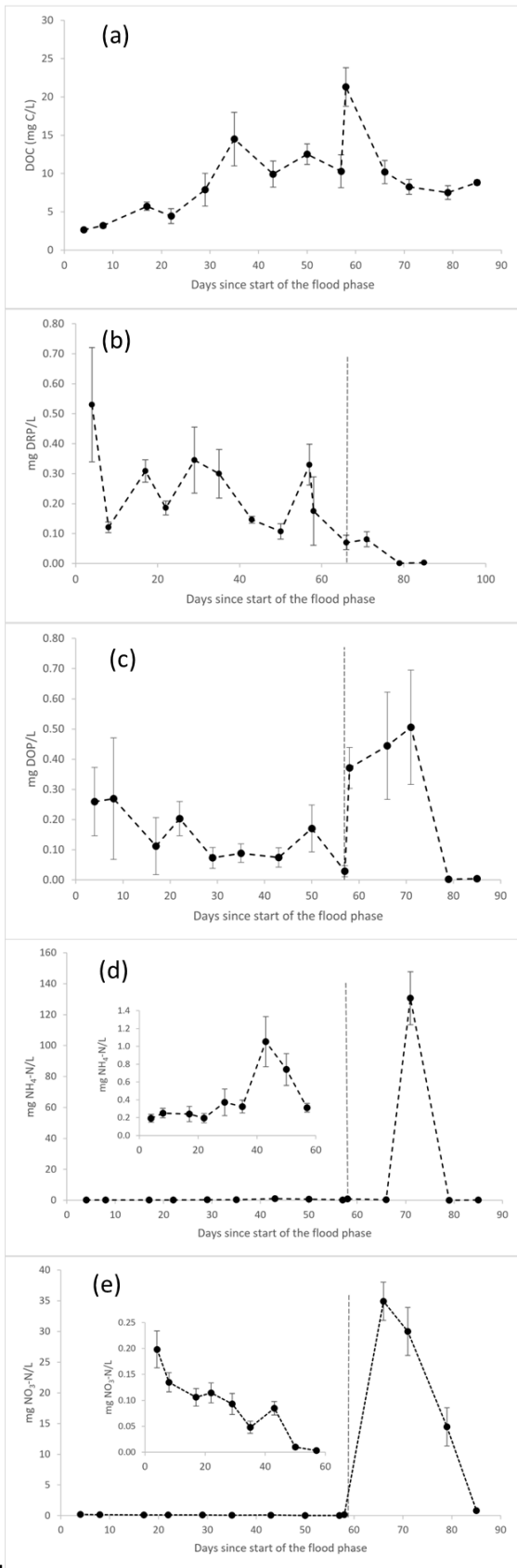


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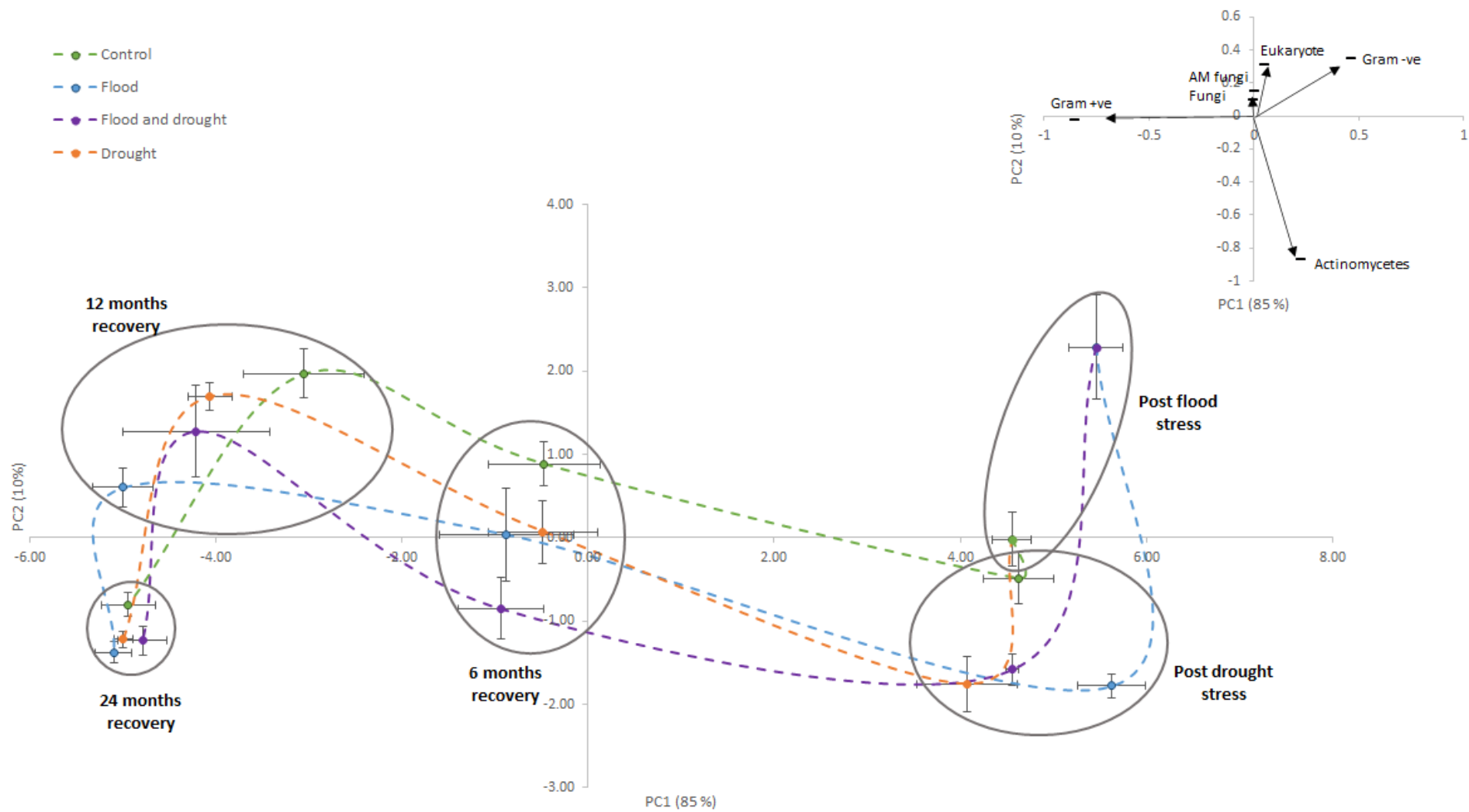
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Figure 2



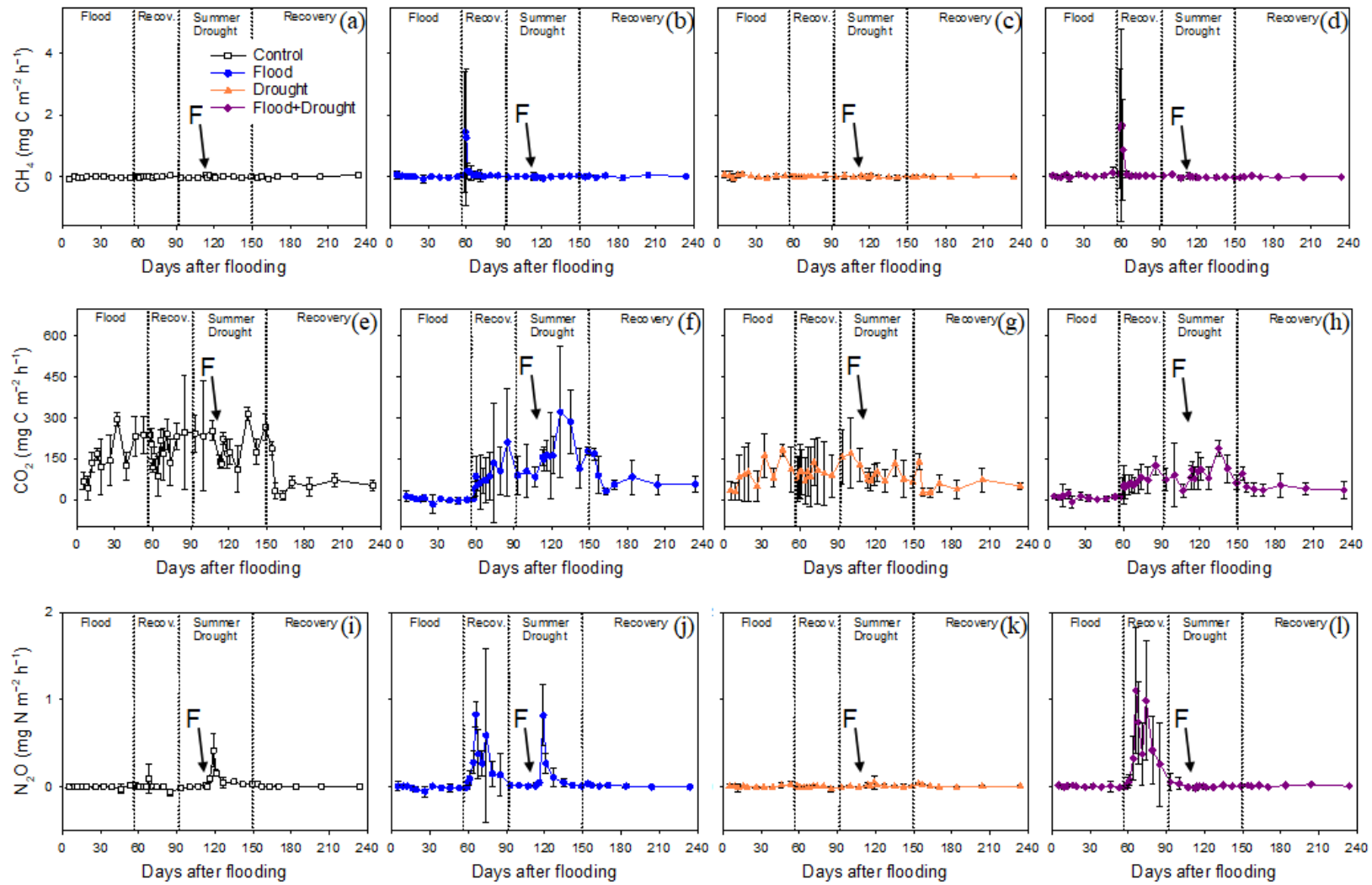
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Figure 3



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Figure 4



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Figure 5